



1 **Single-particle Raman spectroscopy for studying**
2 **physical and chemical processes of atmospheric particles**

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11 **Abstract.** Atmospheric particles experience various physical and chemical processes and change their
12 properties during their lifetime. Most studies on atmospheric particles, both in laboratory and field
13 measurements, rely on analyzing an ensemble of particles. Because of different mixing state of individual
14 particles, only average properties can be obtained from studies using ensembles of particles. To better
15 understand the fate and environmental impacts of atmospheric particles, investigations on their properties
16 and processes at a single-particle level are valuable. Among a wealth of analytic techniques, single-
17 particle Raman spectroscopy provides an unambiguous characterization of individual particles under
18 atmospheric pressure in a non-destructive and in-situ manner. This paper comprehensively reviews the
19 application of such a technique in the studies of atmospheric particles, including particle hygroscopicity,
20 phase transition and separation, and solute-water interactions, particle pH, and multiphase reactions.
21 Investigations on enhanced Raman spectroscopy and bioaerosols on a single-particle basis are also
22 reviewed. For each application, we describe the principle and representative examples of studies. Finally,
23 we present our views on future directions on both technique development and further applications of
24 single-particle Raman spectroscopy in studying atmospheric particles.

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32 **1 Background**

33 Atmospheric particles or aerosols have considerable effects on climate and human health.(Seinfeld and
34 Pandis, 2016) In general, atmospheric particles can originate from a wide variety of anthropogenic and
35 natural sources. Primary particles are emitted directly as a liquid or solid (e.g., soot particles, mineral
36 dust), and they can also undergo atmospheric multiphase reactions. Secondary particles can also be
37 generated by gas-to-particle conversions, such as new particle formation via nucleation and
38 condensation.(Pöschl, 2005) During their atmospheric lifetime, both primary and secondary particles are
39 subject to physical and chemical processes, such as partitioning and multiphase reactions. For liquid
40 particles, gas-particle partitioning takes place to achieve thermodynamic equilibrium. For example, an
41 increase or decrease in relative humidity (RH) would result in the uptake of the water vapor into
42 deliquescent particles or the evaporation of particulate water to the gas phase, respectively. Multiphase
43 reactions can be described as gas-particle partitioning accompanied by chemical reactions. Furthermore,
44 photochemistry is an important group of multiphase reactions during the daytime. Both partitioning and
45 chemical reactions can change the composition of the particles, altering their properties such as
46 hygroscopicity, cloud condensation nuclei (CCN) activities, phase state, morphology, reactivity, toxicity,
47 and finally, environmental impacts.

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49 Physical and chemical processes of atmospheric particles have been examined by numerous experimental
50 approaches: a flowing stream of aerosol particles in an open system, a chamber of suspending particles,
51 a collection of deposited particles, and single particles. An aerosol flow tube is one of the most common
52 approaches employing a flowing stream of aerosol particles. It enables real-time characterization of
53 reaction products and particulate size under controlled conditions (e.g., RH, temperature), with
54 instruments connected to the exit or different outlets along the flow tube reactor. Oxidation flow reactors
55 (OFR) are analogous to the general aerosol flow tube, but photooxidants (e.g., OH radical) are introduced
56 to examine secondary organic matter formation.(George and Abbatt, 2010;Keller and Burtscher,
57 2012;Kang et al., 2007;Smith et al., 2009) Typically, higher concentrations of reactive precursor gases
58 than ambient concentrations are used in these flow systems to compensate for the relatively short
59 residential time (e.g., seconds) of the reactants. However, such equivalency of exposure may not
60 necessarily hold.(Lee and Chan, 2007b;Chu et al., 2019) The smog chamber enables simulations of
61 various environmental conditions, including gas concentrations, for studying the formation and chemical



62 transformation of particles. The residence time of particles can be extended to a few hours, but other
63 limitations exist, such as long response time to control/change conditions, high cost, and wall loss of
64 particles.(Kang et al., 2007) Deposited particles can be studied using a sample-loaded substrate/filter
65 placed in a flowing-gas reactor or a flow tube with its internal wall coated with particles. Environmental
66 conditions can be easily varied, and the reaction time can be very long. An in-situ analyzer for
67 spectroscopic analysis can be equipped (e.g., Raman, Fourier Transform Infrared Spectrometry (FT-IR))
68 since the particles are immobilized on a substrate/filter. However, using a large number of particles or
69 large particles is usually required to achieve sensitive detection. The influence of particle loading on the
70 reactions and chemical characterization is a challenging question. In general, these approaches produce
71 statistical results of a collection of particles. They may not truly reflect the interaction between gas and
72 particle since the composition of the individual particles may not be identical.(Ault and Axson, 2017)

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74 Single-particle characterization has the advantage that the physical and chemical processes undertaken
75 by the particle can be unambiguously monitored. While there is a wealth of instruments for
76 comprehensively characterizing single particles, some are destructive (e.g., mass spectroscopy) or
77 operate under vacuum conditions (e.g., electron microscopy), making them unsuitable for monitoring the
78 gas-particle interaction on the same particles. Single-particle spectroscopic methods (e.g., Raman) can
79 be good alternatives, but the particles are usually supermicron and up to tens of microns. The findings
80 from particles much larger than ambient ones need to be interpreted carefully for drawing implications
81 in submicron atmospheric particles.

82
83 Despite its requirements of relatively large particles for spectroscopic signals, single-particle Raman
84 spectroscopy offers several advantages for studying the atmospheric process of aerosol particles. First of
85 all, a single particle can be unambiguously characterized *in situ* without destruction. From the Raman
86 spectra, chemical compositions of typical aerosol particles (e.g., soot, sulfate, organics, etc.) can be
87 identified by their characteristic peaks. Interactions between molecules can be inferred by peak shifts
88 and changes in peak width. Particulate properties such as the phase state and size can also be
89 characterized by coupling levitation (or isolation) technique and optical microscopy. We will elaborate
90 on the applications of these functions later.

91



92 Furthermore, Raman analysis is operated under atmospheric pressure and enables the use of atmospheric
93 relevant concentrations for a long exposure time (Lee and Chan, 2007a) to examine atmospheric gas-
94 particle interactions. Micro-Raman systems provide the spatial information of chemical species at 2-D
95 or even 3-D levels. (Tripathi et al., 2009; Batonneau et al., 2006; Offroy et al., 2015; Sobanska et al.,
96 2014; Sobanska et al., 2015; Sobanska et al., 2006; Falgayrac et al., 2018; Scolaro et al., 2009b; Ao et al.,
97 2020) Recent advances in stimulated Raman scattering by micro-Raman further improves the spatial
98 resolution from 1 μm to 350 nm and significantly shortens the scanning time from 2.5 h to 2 s. (Ao et al.,
99 2020) The disadvantage of Raman spectroscopy is that it can only detect Raman-active functional groups
100 but cannot provide information on the exact molecular structures. Sometimes, light-absorbing chemicals
101 such as HULIS or some brown carbon species may generate fluorescence, interfering with the peak
102 identification in the spectra. Therefore, single-particle Raman spectroscopy is more suited for laboratory-
103 generated particles than ambient particles. FTIR does not suffer from fluorescence since it uses a longer
104 wavelength (up to tens of μm), but the long wavelength also lowers the spectral resolution. (Ault and
105 Axson, 2017; Liu et al., 2020) Also, Raman has advantages over FTIR for water-containing samples
106 because the strong water peaks in the latter can mask the peaks from other components or even make
107 them undetectable. (Nishikida et al., 2018)

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109 This paper will introduce the general setup of the single-particle Raman spectrometer, how single
110 particles can be characterized unambiguously, and examples of physical and chemical processes reported
111 in the literature. We focus on applications of atmospheric relevance and refer readers to other literature
112 on the principles of Raman spectroscopy and related techniques (e.g., stimulated Raman or coherent
113 Anti-stokes Raman spectroscopy) and applications with less direct atmospheric relevance (e.g.,
114 pharmaceuticals). (Tolles et al., 1977; Jones et al., 2019; Efremov et al., 2008) Finally, we will provide
115 suggestions on future directions of Raman spectroscopy for atmospheric particle studies.

116

117 **2 General set up of single-particle Raman spectroscopy**

118 **2.1 Particle isolation**

119 In general, single-particle Raman spectroscopy consists of two key components: particle isolation and
120 Raman spectroscopy, including micro-Raman and confocal Raman configurations. One of the most
121 popular approaches for single-particle isolation is particle levitation, which has been discussed in a few



122 reviews.(Reid and Mitchem, 2006;Mitchem and Reid, 2008;Krieger et al., 2012) In this paper, we will
123 briefly introduce their principles and configurations and focus on atmospheric applications. Techniques
124 for levitation of particles can be mainly divided into electrical, optical, and acoustic trapping. Another
125 approach for particle isolation is particle deposition onto substrates with Raman analyses of single
126 particles on an individual basis.

127

128 **2.1.1 Electrodynamic balance**

129 The Electrodynamic balance (EDB) has been used for studying single particles with a diameter of 1-
130 100 μm since the 1980s (Davis and Ray, 1980). A particle is levitated and trapped at the null point under
131 the alternating current (AC) and direct current (DC) electric fields in the EDB. In the absence of other
132 forces, the gravitational force of the particle trapped in the EDB is equal to the balancing electrostatic
133 force, as shown in Eq. (1)

$$134 \quad mg = nqC \frac{V_{DC}}{z} \quad (1)$$

135 where m is the mass of the particle, g is the gravitational constant, n is the number of elementary
136 charges present on the particle, q is the elementary charge, z is the distance between two electrodes
137 for the DC voltage, C is a constant dependent on the geometrical configuration of the EDB, and V_{DC}
138 is the balancing DC voltage required to levitate the particle stationary. According to Eq. (1), the mass of
139 the particle is proportional to the DC voltage needed to balance its gravitational force, assuming that the
140 charge present on the trapped particle remains constant. Hence, the relative mass change in the particle
141 undergoing any physical or chemical processing can be quantified by measuring the balancing DC
142 voltage. With a particle levitated stationary, Raman spectroscopy of the stably trapped particles can
143 provide chemical information. Coupling with mass change measurements and optical microscopy,
144 physical and chemical characterization of single particles can be in-situ monitored.

145

146 **2.1.2 Optical trapping**

147 Optical trapping utilizes the radiation pressure from a light source to balance the weight of the levitated
148 particle.(Ashkin, 1970) During optical trapping, the particle interacts with a highly focused laser beam
149 and experiences scattering and gradient forces.(Mitchem and Reid, 2008;Ashkin and Dziedzic, 1981)
150 The scattering force is proportional to the square of the induced dipoles and the intensity of the incident



151 laser beam, acting as a force along the direction of the beam propagation. In contrast, the gradient force
152 derives from the focusing of the laser beam and draws the particle towards the region with the highest
153 light intensity.(Krieger et al., 2012)

154

155 Optical levitation and optical tweezer are two typical optical-trapping techniques utilizing these two
156 forces. While optical levitation relies on the exact balance between downward gravitational force and
157 upward scattering force, optical tweezers effectively create a strong intensity gradient in three dimensions
158 by tightening the focus of the laser beam to generate amplified gradient forces. Optical tweezers provide
159 a more robust trap with the gradient force typically many orders of magnitude larger than the scattering
160 and gravitational forces, leading to 3-dimensional confinement in position with a single laser beam.
161 However, particle size changes must be accompanied by dynamic changes in the light intensity for
162 retaining the trapped particle. In principle, transparent non-absorbing particles of 1–100 μm in diameter
163 can be trapped with dynamic light intensity adjustment for droplet size changes. By configuring the
164 Bessel beam in either vertically or horizontally trapping arrangements, optical levitation and tweezer can
165 be applied for trapping submicron particles(Meresman et al., 2009). However, Raman spectroscopy
166 would require larger particles for sufficient signal-to-noise ratios. A recent study reported a
167 photochemical study of aerosol particles using an optical trap for the first time by employing a custom-
168 built photochemical levitation chamber.(Gómez Castaño et al., 2019) A sidewall of the chamber is
169 equipped with quartz windows for a UV lamp.

170

171 **2.1.3 Acoustic levitation**

172 Compared to the EDB and optical trap, acoustic levitation is less utilized in atmospheric research due to
173 the complicated acoustic streaming and the requirement for large particles (20 μm - 2 mm).(Cohen et al.,
174 2020;Jones et al., 2021) In an acoustic levitator, high-frequency sound waves generated by a piezo-
175 electric radiator reflect from a concave-shaped reflector to generate a standing wave. The latter produces
176 acoustic radiation pressure on the particle, levitating and trapping it against the gravitational force. It has
177 been reported that a single particle can be firmly trapped by acoustic levitation under 15 h continuous
178 irradiation.(Tobon et al., 2017;Seng et al., 2018)

179

180 **2.1.4 Deposited particles**



181 In addition to single-particle levitation, deposited particles on a surface provide a convenient platform
182 for single-particle Raman spectroscopy. While levitation can trap few particles for investigation, surface-
183 based isolation affords a relatively large number of single particles, which can support further off-line
184 characterizations. However, the direct contact between particle and substrate may induce solute
185 nucleation or phase separation that causes unexpected modification in particulate morphology. Hence,
186 hydrophobic substrates are often used to study supersaturated aqueous droplets by reducing the particle
187 and substrate contact area. Investigations on photochemistry can be easily performed by coupling a
188 transparent substrate and an additional light source beneath the quartz window at a flow cell's
189 bottom.(Gen et al., 2019b;Gen et al., 2020) A single particle can be retained for an infinite time deposited,
190 facilitating studies of slow processes in the atmosphere. However, Raman scattering from the substrate
191 could also show interference.

192

193 **2.2 Raman spectroscopy**

194 A single particle Raman spectroscopy set-up generally involves a laser for excitation, followed by
195 collection optics to detect the Raman signals analyzed by a spectrometer with a monochromator and a
196 CCD detector. The Raman peak location and intensity/area provide information on the speciation and
197 abundance of specific particulate components. Appendix A summarizes the Raman peak assignment of
198 representative atmospheric species. Raman peak can also shift its location (in wavenumber) and change
199 in the full width at half-maxima (FWHM) as a response to variations of local bonding environments,
200 such as molecular interaction and phase transition. Furthermore, Raman emissions from spherical
201 particles (e.g., levitated droplets) can be complicated by resonance modes at discrete wavelengths, or so-
202 called size parameters in the spectra. These resonant modes are often called morphology-dependent
203 resonance (MDRs) or whispering gallery modes (WGMs). They can accurately determine size with a
204 resolution down to ~5 nm (corresponding to an uncertainty of ~0.058%)³⁹, refractive index, and even
205 shell-core morphology of the mixed organic-inorganic particles^{39, 40} based on the Mie theory. Micro-
206 Raman is a combined system that uses an optical microscope for guiding the laser to analyte particles
207 and collect Raman scattering from the same microscope connected to a spectrometer. The Raman
208 scattering signal-to-noise ratio can be enhanced since the incident laser energy is focused on a small
209 volume through a high magnification (and short working distance) objective. Furthermore, 2D Raman
210 imaging is possible in studying particles in the order of tens of microns.(Eom et al., 2016;Zhao et al.,



211 2018b) Confocal Raman micro-spectroscopy eliminates the Raman scattering from volumes other than
212 the focal volume by configuring additional slits in the spectrometer. Confocal Raman provides depth
213 profiles of constituent information of the particle.(Scolaro et al., 2009a)

214

215 **3 Characterization of single particles**

216 **3.1 Hygroscopicity and phase transition**

217 Depending on the environmental RH, type of solutes, and size of the particles (Kelvin effect),
218 atmospheric particles can absorb/release an appropriate amount of water to reach equilibrium with the
219 gas phase, i.e., the partial water vapor pressure above the particle surface equals to that of the gas phase.

220 Hygroscopic properties include water uptake/evaporation and phase transition as a response to changes
221 in ambient RH. Upon water uptake/evaporation, the particle size change alters the particulate capacity in
222 scattering light, and particles can also be activated to cloud condensation nuclei.(Tang et al., 2019) In
223 this paper, we focus on subsaturated (RH <100%) environments. The phase states determine the water
224 uptake behaviors of the atmospheric particles. They can be generally divided into liquid (i.e., aqueous),
225 crystalline, and amorphous (e.g., ultra-viscous, gel, rubber, glasses). Both crystalline and amorphous
226 particles are solids, but only the latter can absorb water reversibly, while the former shows hysteresis in
227 water uptake. At increasing RH, water uptake increases the aerosol liquid water (ALW) content of
228 deliquescent/amorphous particles and their size and ability to scatter light. A recent study showed that
229 the poor visibility during haze in the North China Plain is positively correlated to the ALW(Wu et al.,
230 2018). When crystalline particles are exposed to increasing RH, they will initially adsorb a small amount
231 of water on its surface, the so-called surface adsorbed water (SAW), and then uptake a significant amount
232 of water vapor abruptly to become aqueous droplets at the deliquescence RH (DRH). The DRH is the
233 RH at which a droplet attains the saturated salt concentration, ignoring the influence of the Kelvin effect,
234 which is often the case for single-particle Raman analysis (particle size > 10 μm). At decreasing RH, a
235 droplet loses water to become supersaturated and finally may effloresce to form crystalline particles at
236 the efflorescence RH (ERH). Typical inorganic atmospheric components of $(\text{NH}_4)_2\text{SO}_4$, NH_4NO_3 and
237 NaCl deliquesce at ~80%, ~62% and ~75% RH but effloresce at 34-48%, 0-30% and 39-54%,
238 respectively. Many studies show that nitrates and some organics do not effloresce even at low RH. (e.g.,
239 below 5%)(Yeung and Chan, 2010;Mikhailov et al., 2009) In general, organics with high solubility (e.g.,
240 malonic acid) tend to remain a supersaturated liquid state or amorphous state at low RH.(Krieger et al.,



241 2012) Unlike DRH, which is thermodynamically well defined, ERH is a stochastic property that relates
242 to the kinetics of nucleation and presence of impurities. It has been reported that sucrose crystals
243 deliquesce at 85.6%RH upon humidification, but the aqueous droplets turn into a glassy amorphous state
244 upon drying instead of forming crystallized solids.(Zobrist et al., 2011) Using EDB and optical trap
245 experiments, hygroscopic changes and phase transition RH values that involve rapid changes in mass or
246 size can be measured by the balancing voltage and light scattering, respectively. The growth factor (GF)
247 in terms of mass or size ratios as a function of water activity or RH, DRH, and ERH can be easily
248 determined by changes in mass/size and the light scattering pattern.(Mitchem et al., 2006;Chan et al.,
249 2005)

250

251 While GF, DRH, and ERH measurements are readily made for mixed particles using EDB and optical
252 traps, the formation of mixed-phase particles before full deliquescence is more complicated. For example,
253 unlike the corresponding single inorganic salts, NaCl-malonic acid and $(\text{NH}_4)_2\text{SO}_4$ -malonic acid mixed
254 particles absorb a significant amount of water before their DRH since malonic acid absorbs water
255 reversibly without crystallization.(Lee et al., 2008a) However, this stepwise water uptake and phase
256 transition behaviors could not be explained by EDB/optical levitation due to the lack of chemical
257 characterization. Raman spectroscopy offers complementary information on the water uptake behaviors
258 and phase states of Raman active species. For example, the water-to- solute ratio (WSR) as a function of
259 ambient RH can be easily determined from the integrated peak area ratio of the water $\nu(\text{O-H})$ to that of
260 the solute (e.g., sulfate, nitrate) after proper calibration. Also, the full width at half maxima (FWHM) of
261 the characteristic peaks is sensitive to the phase state of the solute. In general, a species in the liquid
262 phase has a larger FWHM than in the solid phase because the less compact molecular arrangement allows
263 minor variations in vibration. In Fig. 1a,b, as the RH decreased to 39.5% for ammonium sulfate (AS)
264 and 19% for ammonium nitrate (AN), crystallization can be inferred by the red-shift and decrease in
265 FWHM of sulfate and nitrate peaks, and the disappearance of water peak. On this basis, the mixed-phase
266 particle prior to full deliquescence can be characterized by FWHM of different Raman-active
267 components. For instance, an increased FWHM of nitrate peak was found ($\sim 730 \text{ cm}^{-1}$) at $\sim 67\%$ RH,
268 followed by an increased FWHM of sulfate peak ($\sim 450 \text{ cm}^{-1}$) at $\sim 76 \text{ RH}\%$ for mixed-phase AS-AN
269 particles when RH increased (Fig. 1c,d). These changes indicate that AN deliquesced before AS did upon
270 humidification, and the mixed-phase AS-AN particle was composed mainly of aqueous AN and



271 crystalline AS. The accuracy of single-particle Raman spectroscopy in determining hygroscopicity has
272 been verified by thermodynamic models such as E-AIM and UNIFAC.(Peng et al., 2001;Yeung and Chan,
273 2010)

274

275 **3.2 Solid-phase transition**

276 In addition to liquid-solid phase changes, transitions can occur between solid phases, from a metastable
277 one to a stable one at a specific temperature. Many organics are known to form polymorphs, defined as
278 solids with the same chemical composition but distinctive properties such as thermal stability and
279 solubility due to different structures.(Krieger et al., 2012) Different polymorphs of the same compound
280 can exhibit other Raman characteristics, which can identify different solid states and their transition. For
281 example, crystalline glutaric acid (GA) particles could exist as the metastable α -form and the
282 thermodynamically stable β -form at room temperature, and α -form GA tends to transfer into β -form.
283 Using EDB-Raman, Chan et al. coated GA on AS particles by coagulation of small GA particles formed
284 by nucleation. Raman spectra showed the existence of α -form GA in the freshly coated particles, but
285 then β -form was detected after a deliquescence - crystallization cycle.(Chan et al., 2006) Ling and Chan
286 reported the formation of the α -form GA solid in a freshly crystallized GA-AS particle. The α -form GA
287 solid gradually transformed to the β -form at 60% RH in 3 h at room temperature.(Ling and Chan, 2008)
288 Furthermore, Yeung et al. examined the effects of polymorphism on the hygroscopic properties of GA
289 particles using Micro-Raman(Yeung et al., 2010). By evaporation at room temperature, α -form and β -
290 form solid particles were formed (α : $\beta \approx 5$: 1). It was observed that the α -form deliquesced at 85-86%
291 RH, whereas the β -form deliquesced at 90% RH. Furthermore, the α -form is transformed to the β -
292 form at 86% RH instead of achieving full deliquescence (Fig. 2). This solid phase transition behavior
293 was used to explain the discrepancies of DRH values of glutaric acid particles reported based on different
294 literature measurements.(Cruz and Pandis, 2000;Peng et al., 2001;Zardini et al., 2008;Parsons et al.,
295 2004;Treuel et al., 2009;Choi and Chan, 2002) Polymorphism also exists in atmospheric relevant
296 inorganics. For example, AN can exist in five stable polymorphic forms with distinct Raman spectra,
297 depending on the temperature. In atmospheric research, very little attention has been paid to the solid
298 phase transitions of AN because phase IV AN particle is stable over a wide range of tropospheric
299 temperatures ($-17^{\circ}\text{C}\sim 32^{\circ}\text{C}$).(Wu et al., 2007) Nevertheless, Wu and co-workers reported that the
300 KNO_3/AN mixed particles undergo the IV-III transition under ambient temperatures at a relatively high



301 mass percentage of KNO_3 (≥ 7.4 wt%), or even crystallize directly in phase III from droplets with a further
302 increase in the mass percentage of KNO_3 . (Wu and Chan, 2008)

303

304 Another typical solid phase transition is between stable/metastable double salts. Like
305 metastable/stable polymorphs, metastable double salts show Raman characteristics different from stable
306 ones, likely due to specific molecular interactions. Ling et al. observed the double-salt $3\text{AN}\cdot\text{AS}$, not
307 predicted from thermodynamics, in the freshly crystallized mixed single particles of AN and AS using
308 single-particle Raman spectroscopy. The existence of double salts indicates that AN and AS likely
309 crystallized simultaneously. The degree of metastability could depend upon the crystallization processes
310 of the particles. Besides, it takes several hours to more than a day, depending on RH, for a freshly
311 crystallized metastable AS-AN double salt particle to reach its stable form, suggesting a high possibility
312 that metastable solids persist in the atmosphere when the ambient RH is sufficiently low. The AN-AS
313 double salts were recently reported in ambient particles during haze events in Beijing with a considerable
314 amount. (Sun et al., 2019b)

315

316 **3.3 Molecular interactions between particulate water and solutes**

317 In general, characterizations on phase transitions relies on the sensitivity of Raman spectroscopy in
318 probing the bonding environment of components in the particles. Raman can also characterize the
319 interactions between the solutes and particulate water, such as hydrogen/ionic bonding. For example, the
320 Raman shift of strongly hydrogen-bonded water (~ 3500 cm^{-1}) is generally higher than that of water
321 monomers (~ 3300 cm^{-1}). (Zhang and Chan, 2003) At low RH, the elevated concentration of electrolytes
322 leads to more significant disruption of the hydrogen bonding environment of the water molecules,
323 resulting in a shift of OH stretching peak at 3480cm^{-1} to higher frequencies. In contrast, the low-
324 frequency shoulder at 3290 cm^{-1} diminishes. Thus, the shape of water peaks in Raman spectra can reflect
325 the ionic strength of the particle, which is an essential parameter for aerosol chemistry. (Mekic et al.,
326 2020; Mitchem et al., 2006) The sensitive shift of water peak has also been used in studying ice nucleation
327 properties such as depositional freezing (Wise et al., 2012; Wise et al., 2010) and immersing
328 freezing (Knopf et al., 2003; Knopf et al., 2002). Specifically, while liquid water peaks are usually located
329 at $\sim 3400\text{cm}^{-1}$ with a shoulder at a lower wavenumber, a sharp peak at ~ 3150 cm^{-1} with shoulder peaks at
330 $\sim 3200\text{cm}^{-1}$ appears as water is frozen. (Mael et al., 2019; Schill et al., 2014) Besides, ionic bonding of



331 inorganic salts (e.g., SO_4^{2-} , NO_3^-) can also be reflected by their Raman characteristics; and many of them
332 have been further supported by theoretical calculations.(Yu et al., 2012;Zhu et al., 2018;Wang et al.,
333 2015b) For example, Zhang et al. observed a blue shift of sulfate and nitrate Raman peaks and an increase
334 in FWHM when the molar water to the solute ratio (WSR) in MgSO_4 and $\text{Mg}(\text{NO}_3)_2$ particles decreases
335 to less than six, the number of hydration of Mg^{2+} ions.(Zhang and Chan, 2000;Zhang et al., 2004) These
336 observations were attributed to direct contacting ion pairs formed between Mg^{2+} and anions, without
337 water molecules between the ions. Besides, single-particle Raman spectroscopic analysis can reveal the
338 structural heterogeneity of droplets. Raman imaging suggested that the NO_3^- to Mg^{2+} ratio of the contact
339 ion pairs increases but the H_2O to Mg^{2+} ratio decreases as the surface of a highly supersaturated droplet
340 is approached.(Zhang et al., 2004) Yeung et al. reported that though no crystallization was shown by
341 glyoxylic acid particle even at 0% RH, a new C=O peak at $\sim 1820\text{cm}^{-1}$ appeared together with the aqueous
342 C=O peak at 1735cm^{-1} , possibly indicating the presence of anhydrate.(Yeung and Chan, 2010)

343

344 **3.4 Liquid-liquid phase separation**

345 As mentioned above, partial crystallization can lead to a separation of solid and liquid mixed phases.
346 Separation of two or even more liquid phases, namely liquid-liquid phase separation (LLPS), is also
347 possible. LLPS can be triggered by the interactions between inorganic, organic solutes and particulate
348 water in mixed particles containing organics and inorganics.(You et al., 2012) This phenomenon is
349 mainly attributed to the low solubility of some non-polar or weakly polar organics, especially in the
350 presence of inorganic salt, which may further decrease the solubilities of organics as the salting-out
351 effects.(Setschenow, 1889) Single-particle (micro-) Raman spectroscopy has been used for probing
352 LLPS via spatial scanning (micro-Raman), optical imaging (micro-Raman), Raman imaging, and
353 whisper gallery modes (WGMs).(Bertram et al., 2011;Song et al., 2012a;Buajareern et al., 2007a) Spatial
354 scanning records distinct micro-Raman spectra at different locations of the particle for indicating
355 LLPS(Bertram et al., 2011). However, this approach requires accurate focusing of the laser onto different
356 targeted phases, which can be difficult when phases are separated into small inclusions.(Buajareern et al.,
357 2007b) Alternatively, optical microscopy can directly indicate LLPS by the contrasting colors on
358 appearance, based on the different refractive indices of materials in different phases. The RH for first
359 observing the separated inclusions in the image is defined as separation RH (SRH).(You et al., 2012;Song
360 et al., 2012b;You et al., 2014;You et al., 2013;Song et al., 2013) Using optical microscopy, SRH of a



361 wide range of mixed particles has been correlated with the O:C ratio, organic to inorganic ratio, particle
362 size, and temperature.(Bertram et al., 2011;You et al., 2014) Another approach for identifying LLPS is
363 Raman imaging of the whole particle. Typical Raman measurements bin the vertical pixels of a CCD to
364 enhance the signal to noise ratio while the horizontal pixels represent wave number after chromatic
365 dispersion by a grating. For this imaging of the whole particle, binning is not used and the CCD detector
366 captures a one-dimensional image of the particle after the monochromator. The vertical pixels register
367 the vertical profile of the Raman scattering (e.g., edge to center to edge) of the particles. Chu et al.
368 reported that the core-shell phase separation in a mixed sucrose-ammonium sulfate particle with a sucrose
369 molar fraction of 50% at <5% RH, but not at 70%RH.(Chu and Chan, 2017a) As shown in Fig. 3a, b, the
370 accumulated C-H signal at two sides indicates the presence of an organic-rich coating and an organic-
371 deficient core. In contrast, the even distribution of the C-H signal in Fig. 3c suggests that the particle is
372 likely homogeneous. Additionally, WGMs have also been utilized to infer the occurrence of LLPS and
373 the specific particle morphologies. For example, different resonance structures can appear in the Raman
374 spectra when core-shell or engulfed phase separation occurs.(Buajareern et al., 2007a, b;Kwamena et al.,
375 2010;Song et al., 2013) Sharps peaks generated by WGMs in spectra can be found if the spatial
376 distribution of particulate species is symmetric, such as those in core-shell separated and homogeneously
377 mixed particles though the number, location, and intensity of WGMs peaks are different. In contrast, the
378 engulfed phase-separated particles do not show any WGMs peaks.

379

380 **3.5 Determination of pH**

381 The pH, defined as the activity of proton in aerosol particles, can also be inferred from Raman spectra
382 based on the intensity of the conjugate acid/base ion pairs. Sulfate/bisulfate has been used most frequently.
383 The pH can be determined by substituting the conjugate ion pair concentrations and its dissociation
384 coefficient into Henderson-Hasselbalch (H-H),(Radić and Prkić, 2012) Debye-Hückel theory, and
385 specific ion interaction theory equations, respectively.(Ciavatta, 1980;Craig et al., 2017;Coddens et al.,
386 2019;Craig et al., 2018;Rindelaub et al., 2016) The latter two are more suitable to supersaturated particles
387 since non-ideality has been considered. Rindelaub et al. reported direct measurements of pH in individual
388 deposited particles according to sulfate/bisulfate peaks in Raman spectra, and its variation with
389 RH.(Rindelaub et al., 2016) Craig et al. extended this approach to a range of atmospherically relevant
390 inorganic and organic acid-base equilibria systems covering a wide range of pH (-1~10).(Craig et al.,



391 2017) To eliminate the artifact of contacts between the particle and the substrate, Coddens et al.
392 determined the pH of trapped particles via titration by coalescence of droplets containing strong
393 acids.(Coddens et al., 2019) Boyer et al. further utilized the WGMs for determining the solute
394 concentration in real-time and achieve uncertainties in pH measurements of picoliter droplets ranging
395 from ± 0.03 to 0.06 pH units.(Boyer et al., 2020)

396

397 **4 Multiphase reactions**

398 In Section 3, applications of characterizing single particles have been reviewed. This section mainly
399 focuses on the "processes" occurring on single particles. Taking water-solute interactions as an example,
400 hygroscopicity describes the ability of solutes to absorb water upon equilibrium at fixed RH. Water
401 uptake is a dynamic process, although it is generally too quick to be a limiting step in most atmospheric
402 applications. However, in some viscous matrixes, the water uptake rate could be reduced by orders of
403 magnitude due to the limited diffusion. Taking advantage of the different Raman characteristics of D₂O
404 and H₂O, Davies and Wilson measured the diffusion coefficients of water molecules in trapped viscous
405 particles by monitoring the isotopic exchange rate using Raman spectroscopy.(Davies and Wilson, 2016)
406 Determination of kinetic parameters of non-reactive processes such as partitioning coefficient, diffusion
407 coefficient, and mass accommodation coefficient has been reviewed in another paper(Krieger et al.,
408 2012). Hence, applications on probing multiphase reactions using single-particle Raman spectroscopy
409 will be the focus below.

410

411 **4.1 Multiphase formation of secondary inorganic aerosols (SIA)**

412 Particulate sulfate and nitrate are two major secondary inorganic aerosols (SIA). They have dominant
413 Raman peaks at ~ 980 cm⁻¹ and ~ 1046 cm⁻¹, respectively. Probing multiphase formation of sulfate by
414 single-particle Raman spectroscopy was first employed by Davis and co-workers for absorption of SO₂
415 by alkali-metal hydroxides particles.(Aardahl and Davis, 1996) In atmospheric applications, particulate
416 sulfate formation has been a debating topic since the currently proposed pathways, including O₃, OH
417 radical, NO₂, H₂O₂, and transition metal ions (TMI), cannot entirely account for the high concentration
418 of sulfate observed in haze events such as those in the North China Plain (NCP).(Cheng et al., 2016)
419 Using single-particle micro-Raman spectroscopy, Chan and his group reported that oxidants including
420 NO₂, N(III), OH radicals, etc., generated from particulate nitrate photolysis could effectively oxidize SO₂



421 to sulfate.(Gen et al., 2019b;Gen et al., 2019c;Zhang et al., 2020;Zheng et al., 2020b) Clear increase of
422 the sulfate peak signal was identified in Raman spectra of UV irradiated nitrate particle (Fig. 4). Later,
423 they also reported sulfate formation by multiphase oxidation of SO₂ during FeCl₃ photolysis.(Gen et al.,
424 2020) The formation of nitrate has also been investigated.

425
426 Zangmeister and Pemberton performed kinetic analysis on the reactions between deposited NaCl
427 particles and gaseous HNO₃ based on the increase of the nitrate peak.(Zangmeister and Pemberton, 2001)
428 Tang et al. coupled optical levitation with Raman spectroscopy to investigate N₂O₅ uptake on SiO₂
429 particle as a function of RH and found that particulate nitrate formation increased as RH increased. They
430 conclude that the water content affects the partitioning of HNO₃ between the gas phase and condensed
431 phases (i.e., adsorbed phase and aqueous phase). More surface adsorbed water layers block the reactive
432 sites for N₂O₅ uptake but promote HNO₃ dissolution. Scolaro et al. studied the three-dimensional
433 evolution of the NaCl(100) surface during its reaction with NO₂ as a function of RH, using confocal
434 Raman imaging combined with non-contact atomic force microscopy (AFM).(Scolaro et al., 2009a) At
435 0% RH, a NaNO₃ monolayer capping the NaCl(100) surface formed. NaNO₃ tetrahedral crystals (<0.5
436 μm) formed at 45% RH, and supermicron NaNO₃ rhombohedral plates were obtained under the H₂O
437 multilayer regime (45%<RH<75%). At RH>80%, both crystalline and aqueous NO₃⁻ were found.
438 Reactions were accelerated in the presence of water vapor due to the formation of reactive HNO₃ by NO₂
439 hydrolysis. Ault et al. reported that individual nascent sea spray aerosol particles showed different
440 reactivity toward HNO₃ and N₂O₅, likely due to compositional heterogeneity. After the same exposure
441 of gaseous reactants, some particles showed intense ν(NO₃⁻) peak, whereas ν(C-H) is still the dominant
442 mode in some other particles.(Ault et al., 2014)

443
444 There are also investigations on the simultaneous formation of particulate sulfate and nitrate. Zhao et al.
445 reported a multi-step reaction mechanism of oxidation of SO₂ by NO₂ on deposited CaCO₃
446 particles.(Zhao et al., 2018a)Further, Yu and Zhao et al. reported that the uptake coefficient of SO₂ on
447 CaCO₃ particles in the presence of O₂/NO₂ mixture was much higher than that without O₂.(Yu et al., 2018)
448 The synergy of NO₂ and O₂, which involved mechanisms such as chain reactions, resulted in much faster
449 sulfate formation than the sum of the reaction rates with NO₂ and with O₂ alone. These reports show the
450 usefulness of single-particle Raman spectroscopy in studying the multiphase formation of particulate



451 sulfate and nitrate. It is worth mentioning that degradation of organo-sulfate has been proposed to
452 produce particulate inorganic sulfate recently (Xu et al., 2020b) and organo-sulfate functional groups are
453 also Raman-active. (Bondy et al., 2018) Raman spectroscopy may find applications in the study of
454 organo-sulfates.

455

456 **4.2 Multiphase formation of SOA**

457 Most of the organic compounds are Raman-active though some of them may not have strong signals.
458 Some generate fluorescence in general orders of magnitude stronger than Raman emissions, masking
459 useful Raman features of particles. The characteristic peaks of a single organic compound are generally
460 more than that of an inorganic one due to its complex molecular vibration modes, and the reaction
461 products may further complicate spectra analysis. Therefore, investigation on the multiphase formation
462 of SOA by Raman spectroscopy is limited. However, though identifying the molecular nature of the
463 products is difficult, single-particle Raman spectroscopy can provide useful information during
464 secondary organic aerosol (SOA) formation. For example, Lee et al. observed a gradually elevating
465 baseline in Raman spectra during acid-catalyzed octanol uptake by levitated sulfuric acid particles at 10%
466 RH but not at 50% RH. (Lee et al., 2008b) Though the speciation of products could not be resolved, the
467 increasing fluorescence is evident, attributable to the formed SOA in the particles. Using an EDB, Chan
468 and Chan examined the dependence of nonanal on the presence of hydrophilic and hydrophobic organics
469 in acidic mixed organic-sulfuric acid particles. The presence of hydrophobic organic materials (oleic acid
470 and its reaction products with sulfuric acid) enhanced SOA formation by reactive uptake of nonanal,
471 likely through an enhanced dissolution of nonanal in the particles. (Chan and Chan, 2011) In contrast,
472 levoglucosan, a hydrophilic organic, did not increase the uptake coefficient of nonanal. It is suggested
473 that acid-catalyzed reactive uptake should be examined with the explicit consideration of the role of
474 particle-phase organics that are either initially present or accumulated as a reaction product. Using micro-
475 Raman, Olson et al. investigated the uptake of isoprene epoxydiol (IEPOX) by α -pinene SOA/toluene
476 SOA-coated AS particles. (Olson et al., 2019) Before the uptake, both coated AS particles show core-
477 shell morphology. $\nu(\text{N-H})$ and $\nu(\text{SO}_4^{2-})$ were found at $\sim 3200\text{cm}^{-1}$ and $\sim 976\text{cm}^{-1}$ in the particle core, while
478 strong $\nu(\text{C-H})$ at $2800\text{-}3000\text{cm}^{-1}$ and other moderate/weak stretches of C-H were found in the shell region.
479 After the uptake, $\nu(\text{C-H})$ and $\nu(\text{RO-SO}_3)$ were found in the particle core, indicating SOA formation,
480 such as organo-sulfate. No organo-sulfate was detected in shell regions of both types of particles,



481 although decreases of $\nu(\text{C-H})$ peak were apparent. Generally, Raman analysis is useful for kinetic
482 analysis of reactions (e.g., measuring uptake coefficients), but it is challenging to reveal complicated
483 mechanisms due to the limited molecular identification. For better examining the reaction mechanism,
484 deposited particles can be extracted for off-line analysis using mass spectrometry (MS)(Chan et al.,
485 2013;Li et al., 2008), although the simultaneous mass and phase state information would not be available.

486

487 **4.3 Multiphase oxidation of organic aerosol particles**

488 In addition to reactive uptake of SOA precursors, single-particle Raman spectroscopy has also been used
489 to explore the uptake of oxidants by organic aerosol particles. So far, the reported studies are limited to
490 ozonolysis. Raman spectroscopy is well suited in ozonolysis studies because the $\text{C}=\text{C}$ bonds and the
491 peroxide groups are Raman-active, and ozonolysis does not generate strongly fluorescent products.
492 Using single-particle Raman spectroscopy, Lee and Chan performed in-situ observations of the particle
493 mass, hygroscopicity, morphology of the oleic acid(Lee and Chan, 2007a), linolenic acid, and linoleic
494 acid particles(Lee and Chan, 2007b) during ozonolysis. They found that low ozone concentration
495 ($\sim 200\text{--}250$ ppb) with a longer exposure period (20 h) favors autoxidation to form $\text{C}=\text{C}-\text{C}=\text{C}$ groups;
496 but an extremely high ozone concentration (~ 10 ppm) at shorter exposure time (2 h) tends to form
497 carboxylic acids ($\text{C}=\text{O}$), secondary ozonide (O-O), hydroxyalkyl hydroperoxide (O-H), etc.(Qiu et al.,
498 2020;Enami and Colussi, 2017) Chu et al. further examined the interchangeability between ozone
499 concentration and exposure time by studying ozonolysis of linoleic acid particles in a broader range of
500 experimental conditions. They found the interchangeability holds at ozone exposure ≤ 500 ppb h and $[\text{O}_3]$
501 ≤ 100 ppb for the formation of autoxidation products containing conjugated diene structures but does not
502 hold for the decay of parent linoleic acid or at higher $[\text{O}_3]$.(Chu et al., 2019) These results indicate that
503 extrapolating experimental results of the ozonolysis under high $[\text{O}_3]$ conditions to ambient levels should
504 be carried out with caution. Single-particle Raman spectroscopy also enables us to infer the volatilization
505 of some ozonolysis products. For example, Lee and Chan found that the mass of oleic acid particles was
506 observed to decrease with time while that of linoleic acid and linolenic acid increased on the contrary,
507 which could be potentially explained by the different volatility of the ozonolysis products.(Lee and Chan,
508 2007b) King et al. reported an optical tweezers study of ozonolysis of oleic acid and synthetic seawater
509 droplets. They found clear Raman signatures indicating nonanoic acid and nonanal production, followed
510 by their gradual disappearance, which is also likely due to the volatilization.(King et al., 2004)



511 Furthermore, Dennis-Smith et al. used optical tweezers to study the oxidation of oleic acid droplets
512 containing an inorganic seed with a time-resolution of 1 s to resolve the reactive loss of ozone and the
513 evaporative loss of products. The oleic acid ozonolysis was completed in ~3000 s, but the decrease of
514 droplet volume still lasts for an additional ~5000 s due to the slow volatilization of the products. (Dennis-
515 Smith et al., 2012)

516

517 Chan and Chan studied the roles of the phase state and water content in the ozonolysis of maleic acid
518 (MA) and AS mixed particles. Aqueous MA was found to be more reactive toward ozone than crystalline
519 MA. Interestingly, they also found that higher water content favors more efficient ozonolysis, although
520 the concentration of MA is accordingly lower. (Chan and Chan, 2012) Furthermore, ozonolysis of
521 optically trapped bioaerosol particles has also been reported. In general, lipids and proteins show distinct
522 $\nu(\text{C-H})$ peaks at 2855 cm^{-1} and 2928 cm^{-1} , respectively. (Kline and Treado, 1997; Rygula et al.,
523 2013; Czamara et al., 2015) The ratio of these peaks can be thus used for inferring the different
524 susceptibilities of lipids and proteins during ozonolysis oxidation. (Gong et al., 2019) The products
525 derived from oxidations such as carbonyls (C=O ; $\sim 1700\text{ cm}^{-1}$) (Lee and Chan, 2007b) and oligomer of
526 cysteine (C-S ; $\sim 662\text{ cm}^{-1}$) (Kampf et al., 2015; Movasaghi et al., 2007) are also detected.

527

528 There are also studies exploring other organic surrogates of atmospheric aerosol with more complicated
529 structures. For example, King et al. measured the uptake coefficients of ozone on fumarate and benzoate
530 droplets (as a proxy of HULIS-containing aerosol) and α -pinene droplets (a proxy of biogenic organic
531 aerosol). (King et al., 2008) Different reaction products such as carbonate were identified in Raman
532 spectra, though the chemical characterization is not comprehensive due to their volatility and the
533 detection limits. The reaction of α -pinene in the droplet phase was more than one order faster than in
534 bulk aqueous solutions, likely due to more effective diffusion of ozone in the gas phase and more efficient
535 mass accommodation at the air-droplet interface. (King et al., 2008)

536

537 **4.4 Amine/Ammonia reactions**

538 Ammonia and amines are the most abundant alkaline gases in the atmosphere. They play a significant
539 role in new particle formation and buffering aerosol pH by neutralization of acidic particles. (Zhang et
540 al., 2015; Kerminen et al., 2018; Kirkby et al., 2011; Almeida et al., 2013; Zheng et al., 2020a) There are



541 three types of reactions involving ammonia and amines: (1) Neutralization reaction between acidic
542 particles and ammonia/amines; (2) Displacement of ammonia by amines from ammonium salt particles
543 and vice versa from aminium salt particles; (3) Reactions between ammonia/amines and dissolved di-
544 carbonyls or between gaseous di-carbonyls and ammonium/aminiium salt particles, both would result in
545 the formation of colored compounds (i.e., Brown carbon; BrC).

546

547 Based on the strong vibration signal of the C-H bond ($\sim 2900\text{cm}^{-1}$, sharp peak) from amine/aminiium salts
548 and the N-H bond ($\sim 2800\text{-}3300\text{ cm}^{-1}$, broad peak), the uptake of amine and ammonia can be easily studied
549 using single-particle Raman spectroscopy. Sauerwein and Chan explored the role of phase state in
550 simultaneous uptake of ammonia and dimethylamine (DMA) by acidic particles (e.g., sulfuric acid and
551 oxalic acid).(Sauerwein and Chan, 2017) Anhydrous acidic particles were inert and took up DMA and
552 ammonia presumably by adsorption only. The uptake rates of DMA and ammonia by aqueous acidic
553 droplets were found comparable initially, but once the droplets were neutralized, the stronger base DMA
554 displaced some of the particulate ammonium. On the other hand, if crystallization took place during the
555 uptake, DMA uptake was inhibited, but continuous ammonia uptake gradually displaced the particle
556 DMA ions.

557

558 Chan and Chan reported that the exchange reactions of ammonia/amine vapors in aqueous particles were
559 reversible, and the exchange rates of aqueous particles were, in general, higher than those of their
560 corresponding solid counterparts. Compared with crystalline solids (e.g., methylammonium sulfate),
561 amorphous solid (e.g., diethylammonium sulfate) was found with a higher exchange rate, likely due to
562 the less compact molecular structure.(Chan and Chan, 2013) Chu and Chan further examined the effect
563 of a hydrophilic viscous organic surrogate and a hydrophobic organic coating on DMA uptake by AS.
564 Both the viscous materials and the fresh and aged organic coating retarded DMA uptake by mass transfer
565 limitation.(Chu and Chan, 2017a;Chu and Chan, 2017b) The uptake of DMA by crystalized AS particles
566 at 10-30% RH resulted in the deliquescence of the particles, likely due to the presence of surface adsorbed
567 water (SAW) that promoted the dissolution of DMA. Once dimethylaminiium sulfate is formed, it absorbs
568 more water since it is more hygroscopic than AS and further promoted DMA uptake.(Chu and Chan,
569 2017a) As crystalline solids are generally considered inert towards reactions, this finding indicates that
570 SAW on crystals can serve as a medium to facilitate gas uptake.



571

572 BrC has attracted increasing attention due to its atmospheric abundance and capacity in regulating
573 radiative balance.(Laskin et al., 2015) The reactions between dicarbonyls and ammonium/ammonium-
574 containing salt particles can form BrC. Gen et al. investigated the multiphase uptake of glyoxal by
575 ammonium salt-containing particles using Micro-Raman spectroscopy. They found a gradual elevation
576 of the baseline, which was attributed to the formation of a fluorescent BrC (i.e., 2,2'-biimidazole, BI),
577 based on off-line analysis by UV-VIS spectroscopy and fluorescence spectroscopy.(Gen et al., 2018) The
578 reduction in ALW at low RH increases the glyoxal uptake rate via the enhanced "salting-in" effect and
579 the BI formation rate by facilitating dehydration reactions. Mabato et al. examined glyoxal uptake by
580 methylammonium salt-containing particle. Although the "salting-in" effect was also found in
581 methylammonium salt-containing particles, enhanced BI formation was not observed, likely due to the
582 higher hygroscopicity of methylammonium salt (i.e., higher ALW content) than ammonium salt.(Mabato et
583 al., 2019)

584

585 **4.5 Photochemistry of aerosol particles**

586 A particular type of aerosol particle reaction that deserves special attention is photochemical reactions,
587 which can be initiated by irradiation of the particles. Nitrate can be photolyzed under atmospheric
588 irradiation ($\lambda > 290\text{nm}$). Nitrate photolysis proceeds via three channels: one produces NO_2 and OH
589 radicals, another produces NO_2^- and $\text{O}(^3\text{P})$ and the last generates ONOO $^-$. The third channel's quantum
590 yield is considered negligible compared with the former two since ONOO is unstable. It has been
591 reported that NO_2^- and ONOO $^-$ peaks increased, but NO_3^- peaks decreased in the Raman spectra of
592 acoustically levitated nitrate particles irradiated by UV.(Tobon et al., 2017) NO_2 does not have a large
593 absorption cross-section, but it may be hydrolyzed into NO_2^- .(Tobon et al., 2017;Seng et al., 2018;Gómez
594 Castaño et al., 2019) Most nitrate photolysis products are oxidants, so their ability to initiate oxidation
595 reactions has also been explored(Gen et al., 2019b;Gen et al., 2019c;Zhang et al., 2020). The production
596 of sulfate via nitrate photolysis was discussed earlier. Recently, Zhang et al. reported a prominent formate
597 production by glyoxal oxidation during particulate nitrate photolysis, instead of the more commonly
598 found oxalic acids oligomers in the oxidation of glyoxal in bulk solutions.(Zhang et al., 2021)
599 Furthermore, Liang et al. reported nitrate photolysis in particles containing sucrose, which was used as
600 a surrogate for atmospheric viscous organics. They found that the viscosity of the particle did not reduce



601 the nitrate photolysis rate, but it limited the diffusion of gaseous products (e.g., NO_x, NO_y) from nitrate
602 photolysis. At a high sucrose fraction and 30%RH, the production of these gaseous products and the
603 particles' viscosity transformed the droplets into “hollow and enlarged” semi-solid particles.(Liang et al.,
604 2021)

605
606 Size dependence of nitrate photochemistry has been recently reported, based on single-particle Raman
607 spectroscopy. Tobon et al. studied photolysis acoustically levitated nitrate particles (40~80µm) at 254nm
608 UV radiation. No change in Raman spectra for 2 h. In contrast, smaller particles (5~8µm) trapped
609 optically yielded various Raman peaks of nitrate photolysis products.(Gómez Castaño et al., 2019;Tobon
610 et al., 2017) The slow photolysis in the larger nitrate particles was attributed to the self-inhibited surface
611 process. Additionally, an uneven distribution of NO₂⁻ inside the smaller particle was also observed
612 through Raman mapping.(Gómez Castaño et al., 2019) Organics in the aerosol particle may also
613 decompose under light. For example, Farmentier et al. reported photodegradation of optically trapped
614 oleic acid droplets under 532 nm and 660 nm irradiation, likely initiated by photosensitization. However,
615 they cautioned that degradation loss of optically trapped particles might be artifacts. The photo-decay of
616 light-absorbing particulate matters is also of great interest in understanding their fate in the atmosphere.
617 However, there are very limited studies of atmospheric photochemical reactions using Raman
618 spectroscopy, partly because of the need for complementary chemical speciation for detailed mechanistic
619 studies.

620

621 **5 Enhanced Raman spectroscopy**

622 Though Raman has been used to study physical and chemical processes of atmospheric particles, most
623 of the studies used large particles (dozens of microns) and focused on lab-generated particles with a few
624 representative components. That is primarily because of the intrinsic diffraction limitation of Raman
625 measurements that the laser beam cannot be effectively focused to a spot smaller than its wavelength and
626 thus limits the detection of particles to larger than 1 micron or so. Fluorescence can also be a complication
627 in some chemical systems.

628

629 Various methods have been developed to improve the Raman signal-to-noise ratios and the spatial
630 resolution in other research fields such as material science and biology. Surface-enhanced Raman



631 spectroscopy (SERS) and Tip-enhanced Raman spectroscopy (TERS) have been recently employed in
632 atmospheric research. SERS and TERS enhance Raman emissions by the localized surface plasmon
633 resonances using noble metal nanoparticles (NPs) to provide "hot spots" near the contacting sites between
634 the analyte and NPs.(Ault and Axson, 2017) According to the SERS electromagnetic (EM) enhancement
635 theory, the relationship between the SERS intensity (I) and the local EM strength (E) follows $I \propto |E|^4$.
636 The relationship between E and the distance between NPs and analytes (D) is described as $E \propto$
637 $(1/D)$.(Zhou et al., 2015) Hence, maximizing E and minimizing D without inducing significant
638 perturbation to the sample is the primary goal. In addition to this physical enhancement based on the
639 electromagnetic field, Raman signals can also be enhanced chemically. Specifically, molecular/ionic
640 species can attach to the surface of NPs, distorting the molecular structure and increasing the absorption
641 cross-section. Consequently, the Raman signal of an adsorbed species can be enhanced, potentially
642 accompanied by a peak shift.(Gen and Chan, 2017;Craig et al., 2015)

643
644 SERS enables the measurements of single particles smaller than 1 micron. Craig et al. found that the
645 signals from individual submicron (down to 0.6 μm) particles could be effectively detected with intensity
646 enhancement.(Craig et al., 2015) Furthermore, Tirella et al. extended the SERS detection to the
647 accumulation mode (800~150 nm) particles(Tirella et al., 2018). Sun et al. reported a size-resolved
648 (covering both coarse mode and accumulation mode) SERS characterization study of atmospheric
649 particles sampled during haze events.(Sun et al., 2019a;Sun et al., 2019b) More recently, Kunihisa et al.
650 reported sensitive SERS measurement of Aiken mode particles (<100nm) by combining a condensational
651 growth tube sampler (CGT) and a SERS substrate for directly sampling.(Kunihisa et al., 2020) The wide
652 analytical range of particle size enables SERS to be a potentially powerful tool in studying ambient
653 particulate matters.(Vejpongsa et al., 2017;Steer et al., 2016;Chen et al., 2021)

654
655 Some species including components of microorganisms(Tahir et al., 2020), microplastics(Xu et al.,
656 2020a), crystalline silica(Zheng et al., 2018), the fine structure of organics(Craig et al., 2015;Gen et al.,
657 2018;Fu et al., 2017), and even some Raman-active inorganic salts at very low concentration(Saniel et
658 al., 2019) are invisible in normal Raman due to their weak Raman-activities. However, they could be
659 detected with a SERS substrate. The sensitive detection by SERS can provide valuable insight into
660 particle transformations and emission sources.(Vejpongsa et al., 2017;Ghosal and Wall, 2019;Lee et al.,



661 2019) However, the enhancement factor of SERS is not constant. It strongly depends on the distribution
662 and quantity of noble metal NPs and the hotspots over the substrate, making quantitative analysis difficult.
663 Many attempts have been made to explore preferable and tunable substrate for generalizable atmospheric
664 particle detection,(Cheng et al., 2021;Xu et al., 2020a) but still more progresses are needed in the future.
665
666 The most common SERS application uses particles deposited on SERS substrates, relying on the "hot
667 spot" locating beneath the analyte particles. Recently, some studies create innovative interactions
668 between particulate analytes and SERS NPs to enhance SERS detections. Wei et al. used 4-
669 mercaptobenzoic acid-functionalized AuNPs as pH nanoprobes and internal SERS substrate and found
670 the existence of a stable pH gradient inside aqueous particles.(Wei et al., 2018a) Hence, the assumption
671 of homogeneous mixing might not always be correct, and pH predictions by thermodynamic model might
672 underestimate the acidity of interfacial layers of particles, which is crucial for multiphase uptake or
673 surface reactions. As a further investigation of the SERS pH nanoprobe, Huang et al. reported that the
674 pH distribution inside aerosol particles is affected by the internal distribution of ions, which may depend
675 on the ionic speciation. While sodium tends to reside in the bulk of the particle, leading to a high centroid
676 pH, ammonium distributes more evenly due to its increased surface propensity.(Huang et al., 2020)
677
678 Gen et al. developed an electrospray system (ES-SERS) to deposit SERS NPs onto the surface of the
679 analyte particle and found that this technique can selectively enhance the peak intensity of the species
680 located at surface layers such as surface adsorbed water(Gen and Chan, 2017) and organic coatings(Gen
681 et al., 2019a). In a typical micro-Raman set-up, the objective is above the analyte particle. When the
682 analyte particle is deposited on the SERS substrate, the SERS hot spots are located beneath the particle.
683 In other words, upon laser irradiation from the top, the photons generated by excitation need to pass
684 through the particle with a refractive index greater than air, which would scatter or absorb light and thus
685 reduces the signals received by the Raman microscope. In contrast, ES-SERS hot spots are created on
686 the surface of the particle using electrospray. This configuration reduces the loss of photons and affords
687 greater enhancement in the Raman intensity. (Fig. 5) Besides, as the SERS effect is much more prominent
688 near the hot spots (i.e., small D), the location of the hot spots determines where will be probed in SERS
689 spectra. Therefore, while typical SERS substrate collects enhanced Raman signals from the substrate-
690 particle interface, the ES-SERS can provide information from the gas-particle interface, where many



691 atmospheric multiphase processes occur. Sivaprakasam et al. reported real-time SERS measurements of
692 electrostatically trapped particles containing analyte molecules and NPs as internally mixed liquid
693 particles or coating on solid particles.(Sivaprakasam et al., 2017;Sivaprakasam and Hart, 2021) The
694 interaction between NPs and analyte increased due to more contact areas than that between an aerosol
695 particle and a substrate.

696

697 Similarly, TERS can produce enhanced Raman signals even from particles in the nucleation mode. TERS
698 is a combination of atomic force microscope (AFM) and Raman spectrometer. In general, AFM
699 determines the topography of a single deposited particle by detecting the interaction force between the
700 AFM tip and the surface of the analytes. In a typical TERS set-up, the AFM tip is coated with noble
701 metal particles. An enhanced Raman signal is generated when the tip is interacting with analytes under
702 laser excitation.(Rodriguez et al., 2015) The lack of chemical information of AFM is complemented by
703 SERS. Thus, TERS enables simultaneous characterization of particulate topography and surface species
704 of submicron particles at a high spatial resolution.(Ofner et al., 2016) TERS also has the potential for
705 quantitative analysis since the quantity and configuration of hotspots on the tip are constant.

706

707 **6 Bioaerosols**

708 Bioaerosols have attracted increasing attention worldwide dramatically due to the COVID-19 epidemic.
709 Studies using normal Raman/SERS for off-line detection of biological samples(Félix-Rivera and
710 Hernández-Rivera, 2012;Mosier-Boss, 2017), real-time sensing of bioaerosol(Huffman et al., 2020) have
711 been reviewed in detail elsewhere. Bioaerosols have various Raman active components such as
712 phosphate lipid (δ (=C-H), $\sim 1268\text{cm}^{-1}$; ν (C-H), $\sim 2855\text{cm}^{-1}$), protein (amides bands, $\sim 1260\text{cm}^{-1}$,
713 $\sim 1600\text{cm}^{-1}$, $\sim 1660\text{cm}^{-1}$; C-C backbone stretching in β -sheet proteins, $\sim 983\text{cm}^{-1}$; ν (C-H), $\sim 2928\text{cm}^{-1}$) and
714 nucleic acids (e.g., adenine, $\sim 735\text{cm}^{-1}$), etc.(Socrates, 2004;Talari et al., 2015;Rygula et al., 2013)
715 However, Raman signals from bioaerosol particles are generally weak. To amplify the signals, bioaerosol
716 particles are trapped or deposited onto a substrate for longer integration time.(Tripathi et al., 2009;Laucks
717 et al., 2000;Sengupta et al., 2005;Wang et al., 2015a;Ai et al., 2020;Redding et al., 2015) Using a higher
718 energy laser can also enhance Raman intensity, but it may cause damage to bioaerosol (e.g.,
719 fragmentation).(Gong et al., 2017) On this basis, SERS is more suitable than normal Raman spectroscopy.
720 SERS can also quench the fluorescence generated by biomolecules such as tryptophan and



721 tyrosine.(Gong et al., 2017)

722

723 In general, there are two strategies of single-particle SERS of bioaerosols: (1) directly sampling
724 bioaerosols on a SERS substrate; (2) moving mixed colloids of SERS NPs and bioaerosols onto a glass
725 substrate, followed by drying. While the former is more convenient without complex pretreatments(Tahir
726 et al., 2020), the latter enables flexible modification of NPs properties. For example, functionalized
727 surface of NPs with tag molecules (e.g., antibody), so-called label-based SERS(Liu et al., 2010), were
728 used for detecting specific biomaterials with high reproducibility. Label-free SERS refers to that without
729 tag molecules, such as Au@Ag. In general, by simply mixing, NPs distribute on biomaterials without
730 specific interactions, and the SERS reproducibility is limited. However, by coating AgNPs on the cell
731 wall of bacteria,(Zhou et al., 2014a;Zhou et al., 2015) SERS enhancement was ~30-folds reproducibly
732 higher than simply mixed colloid–bacterial suspension. Furthermore, the discrimination of live and dead
733 bacteria was possible. NPs can be directly formed inside some specific bacteria via bio-reduction for
734 probing the intercellular structure.(Jarvis et al., 2008) Overall, SERS can be an innovative tool in
735 bioaerosol characterization, but the antibacterial effect of NPs and heat effect due to excitation laser
736 should be considered.(Mosier-Boss, 2017) More examples of bioaerosol detection can be found in a
737 recent review.(Liu et al., 2017)

738

739 **7 Future directions**

740 This review paper has shown examples of single-particle Raman spectroscopy in studying atmospheric
741 processes. Overall, the water content, phase state, morphology, and chemical composition of single
742 particles during physical and chemical processes at controlled RH can be monitored in situ.
743 Hygroscopicity and phase behaviors covering a wide range of RH after these processes can also be
744 comprehensively assessed by performing evaporation-humidification cycle measurements.(Chan et al.,
745 2006;Chan and Chan, 2007;Chen et al., 2020;Liu et al., 2008) Although fluorescence may exist in some
746 particles, it could be potentially reduced or even eliminated by (1) using longer excitation wavelengths
747 (e.g., 785 and 1064 nm); (2) employing SERS/TERS; (3) photo-bleaching of fluorescent substances
748 under UV-VIS illumination.(Gong et al., 2017) There are many other innovative studies built on trapping,
749 manipulating, and observing multi-particles via Raman(Davies, 2019;Mitchem and Reid, 2008), such as
750 measuring viscosity and diffusion coefficient(Power et al., 2013), probing reaction(Aardahl,



751 1998;Rkiouak et al., 2014) and phase transition(Richards et al., 2020). Based on these functions and
752 research progresses, we would like to provide the following suggestions on future directions:

753

754 ● **Hyphenated single-particle Raman spectroscopy.** Raman spectroscopy is non-destructive, so
755 incorporating other spectroscopic techniques for characterizing the same particle is possible.
756 Vibrational sum frequency generation spectroscopy (SFG), Fourier transform infrared spectroscopy
757 (FT-IR), and fluorescence spectroscopy, which are also non-destructive, could be complementary
758 to Raman in probing chemical composition of single particle. Recently, novel combination of EDB
759 and mass spectrometry was reported, which can directly introduce the trapped particle for mass
760 spectrometry analysis.(Birdsall et al., 2018;Willis et al., 2020;Jacobs et al., 2017) Adaption of this
761 technique with Raman to form an EDB – Raman system with subsequent MS analysis is relatively
762 straight forward. For optical properties, UV/Vis spectroscopy could be a promising choice.(Jones
763 et al., 2021) While general laser-based MDR can easily bleach light-absorbing molecules, a recent
764 study reported that MDR modes under broad band excitation can measure refractive index (RI) of
765 light-sensitive samples.(Price et al., 2020) Such MDR measurements can complement in-situ Raman
766 spectroscopy for chemical characterization. Additionally, SERS/TERS are emerging useful
767 approaches, but current SERS/TERS methods for atmospheric research mainly use bare NPs. The
768 direct contact of NPs to analyte molecules can potentially induce chemical enhancement, which
769 could complicate the identification of the phase transition and functional groups.(Gen and Chan,
770 2017;Craig et al., 2015) To remove the chemical enhancement, future studies on chemically inert
771 NPs such as shell-isolated nanoparticle-enhanced Raman spectroscopy (SHINERS) are
772 suggested.(Li et al., 2010;Anema et al., 2011)

773

774 ● **Multiphase reaction studies based on single-particle Raman spectroscopy.** Single-particle
775 Raman spectroscopy enables investigations of various atmospheric reactions, including those in
776 surface adsorbed water for solids, mixed-phase particles (solid, aqueous, organic), supersaturated
777 droplets, atmospheric aging of bioaerosols and ultrafast reactions. For solid particles (e.g.,
778 crystalline, mineral dust), multiphase reactions can be “mediated” by surface adsorbed water layers
779 and may result in phase transition, turning the particle into liquid for further reactions.(Kuwata and
780 Martin, 2012;Chu and Chan, 2017a;Zhao et al., 2018a;Yu et al., 2018;Gen et al., 2019a) Probing



781 heterogenous reactions on the particle surface is possible via ES-SERS, TERS, and normal Raman
782 imaging. As particle constituents can be unevenly distributed (e.g., in LLPS) within a single particle,
783 chemistry may not be the same between the bulk and the surface layers.(Wei et al., 2018b;Huang et
784 al., 2020;Olson et al., 2019;Lam et al., 2021;Qiu et al., 2019) The studies of the atmospheric
785 processing of pre-existing particles to form secondary aerosols can be further extended to the aging
786 of bioaerosols.(Gong et al., 2019) Using pulse laser for excitation, time-resolved resonance Raman
787 spectroscopy (TR3) can be a promising tool in probing short-lived reactive intermediates (e.g.,
788 triplet state molecules). In particular, it helps deepen our understanding in ultrafast reactions or
789 initiation steps of typical reactions.(Sahoo et al., 2011;Collins et al., 2018)

790

791 ● **Particle size and temperature effects.** Many publications report single-particle Raman
792 spectroscopy studies under different environmental conditions, including RH and gas concentration,
793 etc. However, how experimental conditions such as temperature and particle size affect kinetics and
794 products in multiphase reactions remains ambiguous. For example, various pathways and kinetic
795 parameters have been proposed to explain particulate sulfate formation in winter haze events in
796 northern China.(Chan and Yao, 2008;Su et al., 2020) However, most studies are based on
797 experimental results at room temperature or 298K. Further exploration with respect to the effect of
798 low temperature on the properties of the particles and multiphase sulfate chemistry is still warranted.
799 Employing a temperature-controlled cold plate flow cell - Raman system for deposited single
800 particles(Roy et al., 2020;Mael et al., 2019;Wheeler et al., 2015), investigations of temperature-
801 dependent reaction kinetics can be easily achieved. Temperature control in a small levitation
802 chamber is also more convenient than in a flow tube or smog chamber. To date, almost all single-
803 particle Raman studies involve supermicron particles. While equilibrium measurements are not
804 sensitive to size (assuming the Kelvin effect is considered), kinetic analysis is often size-dependent.
805 Hence, uptake coefficients for multiphase reactions are usually reported and are restricted to
806 experimental conditions that gas and liquid phase transport are not the rate-limiting step. The size
807 detection limit of single-particle Raman spectroscopy can be extended to sub-10 micrometer with
808 high sensitivity using high-energy excitation laser and tighter optical focusing. The effect of
809 temperature and particle size on multiphase chemistry should be further investigated using long
810 integration time, novel optical configuration (e.g., confocal micro-Raman), and SERS/TERS, etc.



811 Although it is challenging to study single submicron particles by Raman spectroscopy regularly,
812 Raman characterization of a collection of single submicron particles is feasible.

813

814 *Competing interests.* The authors declare that they have no conflict of interest.

815

816 *Acknowledgments.* We gratefully acknowledge the supports from the National Natural Science
817 Foundation of China (No. 41905122, 42075100, and 41875142), the Hong Kong Research Grants
818 Council (No.11302318), the Japan Science and Technology Agency (JST) for Fusion Oriented Research
819 for disruptive Science and Technology program (20352853), the Japan Society for the Promotion of
820 Science (JSPS) Grants-in-Aid for Early-Career Scientists (21K17876).

821

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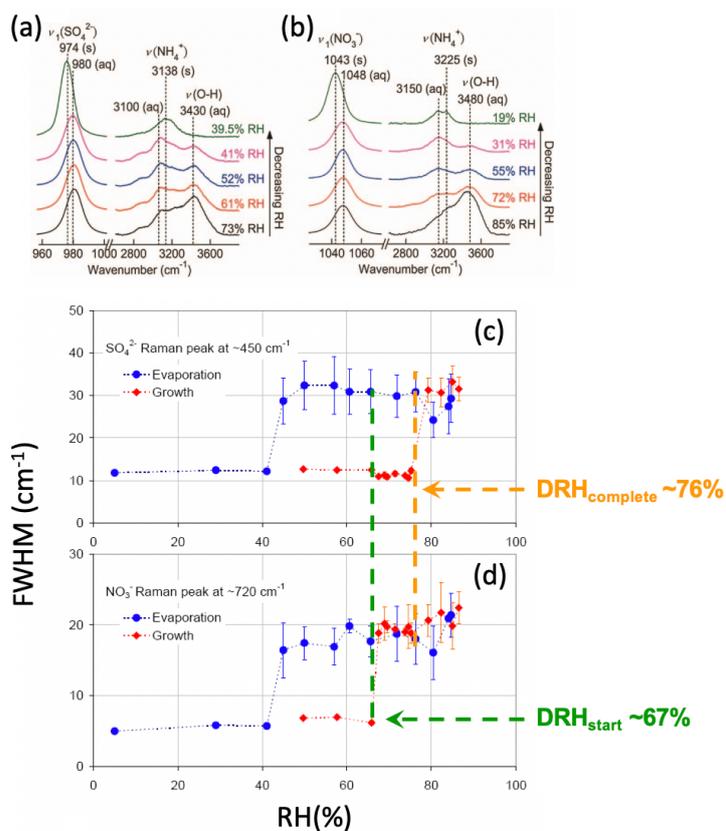


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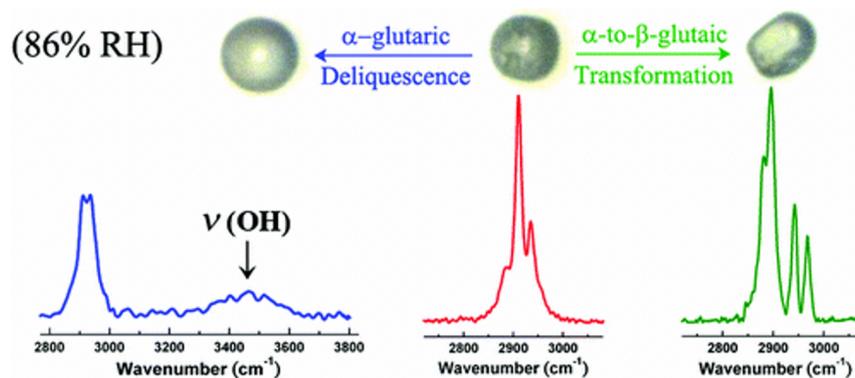


1443

1444 **Figure 1.** Raman spectra of a single (a) AS and (b) AN particle upon evaporation. FWHM of (c) SO_4^{2-}

1445 peak and (d) NO_3^- peak during the evaporation-growth cycle of a mixed AS-AN particle.

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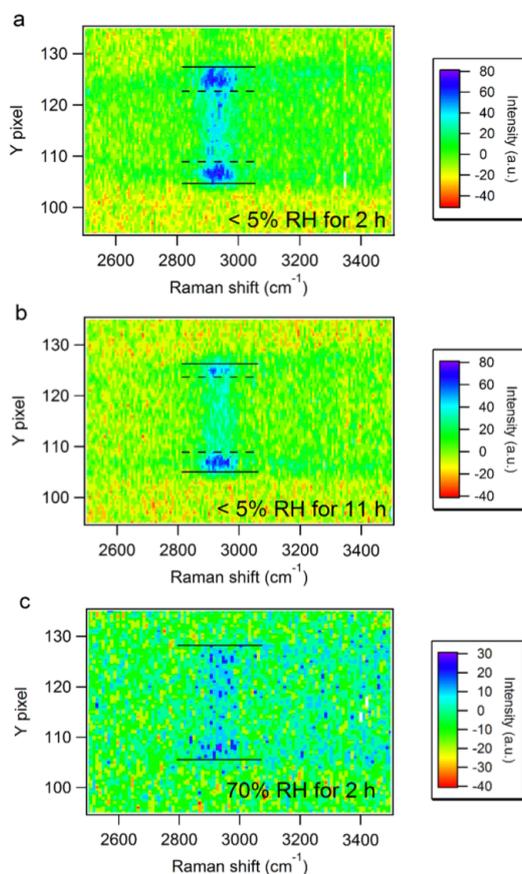
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1448 **Figure 2.** Raman spectra of deliquescent α -glutaric acid, crystalline α -glutaric acid and β -glutaric acid.

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1452 **Figure 3.** Spatially resolved Raman intensities of AS – sucrose (molar fraction of sucrose in total solute

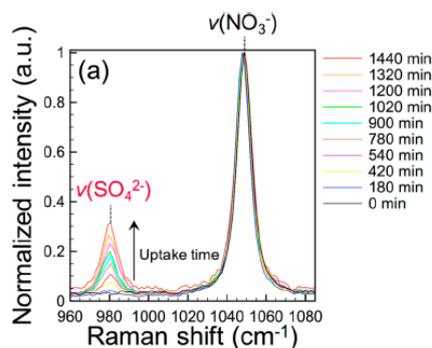
1453 ($F_{\text{SU}} = 0.5$)) particles equilibrated (a) at $< 5\%$ RH for 2 hours and (b) for 11 h and (c) at 70% RH for 2

1454 hours. C-H peaks at ~ 2900 cm^{-1} represent sucrose. Solid lines indicate the particle's upper and lower

1455 edges. The outer regions between the dashed and solid lines are sucrose-rich in (a)-(b). Reprinted with

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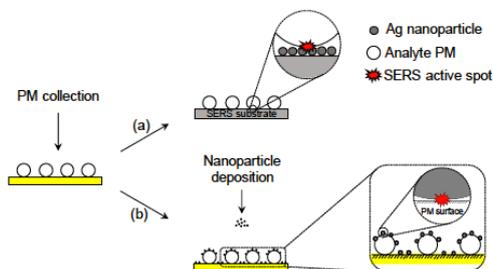


1458

1459 **Figure 4.** Sulfate formation during the photolysis of AN particles in the presence of SO₂ at 80% RH.

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1462

1463 **Figure 5.** Schematic illustration of (a) a conventional SERS substrate approach and (b) the ES-SERS
 1464 approach. Reprinted with permission by Gen and Chan. (2017).

1465

1466 Appendix A. Major Raman peaks of atmospheric relevant species

Species		Raman shift (cm ⁻¹)	Ref.
Water	O-H	~3300 (monomers)	(Carey and Korenowski, 1998)
		~3600 (hydrogen-bonded)	
Inorganic	NH ₄ ⁺	~1413; ~1660; ~3100	(Chu and Chan, 2017a; Sauerwein and Chan, 2017)
	SO ₄ ²⁻	~450; ~618; ~980	(Chu and Chan, 2017a; Ling and



			Chan, 2007)
	HSO ₄ ⁻	~1053	(Rindelaub et al., 2016; Craig et al., 2017; Craig et al., 2018)
	NO ₃ ⁻	~717; ~1046; ~1413	(Gen et al., 2019b; Tobon et al., 2017)
	HNO ₃	~1306	(Rindelaub et al., 2016)
	NO ₂ ⁻	~815; ~1278; ~1330	(Tobon et al., 2017)
	ONOO ⁻	~980	(Tobon et al., 2017; Tsai et al., 1994)
	CO ₃ ²⁻	~1067	(Rindelaub et al., 2016)
	HCO ₃ ⁻	~1019	(Rindelaub et al., 2016)
Soot	D	~1367	(Gen and Chan, 2017; Ivleva et al., 2007)
	G	~1585	(Gen and Chan, 2017; Ivleva et al., 2007)
Mineral dust	SiO ₂	Si-O-Si	~390; ~490; ~795; ~972; ~1460
		O-Si-O	~356
		Si-O	~465
	α-Al ₂ O ₃	Al-O-Al	~420; ~585; ~796;



		Al-OH	~976	2014)	
	Kaolinite	Si-O	~397	(Laskina et al., 2013)	
		Si-O-Al	~516		
		O-H	~914; 3624; 3656; 3668; 3698		
	TiO ₂	Ti-O	~484; ~680;	Amorphous	(Rkiouak et al., 2014;Tang et al., 2014)
			~153; ~639;	Anatase	
		O-O	~250; ~317;	Amorphous	
	CaCO ₃	CO ₃ ²⁻	~713; ~1087	(Laskina et al., 2013;Zhao et al., 2018b;Yu et al., 2018)	
Organic	C-H		~1400; 2700-3000	(Chu and Chan, 2017a;Gen et al., 2018)	
	C-O		~1470	(Laskina et al., 2013)	
	C=O		~1700	(Yeung and Chan, 2010;Yeung et al., 2010;Bondy et al., 2018)	
	O-H		~1700; 3200-3400	(Yeung and Chan, 2010;Zhou et al., 2014b)	
	C-N-C		~890	(Chu and Chan, 2017a)	
	O-C-O (ring)		~770	(Gen et al., 2018)	
	ring		~950	(Gen et al., 2018)	
	O-O		~850	(Lee and Chan,	



			2007a, b)
C=C	~1600		(Lee and Chan, 2007b;Chu et al., 2019)
C=C-C=C	~1640&1655 (double peaks)		(Lee and Chan, 2007b;Chu et al., 2019)
C=C-C=O	~1690		(Chu et al., 2019;Lee and Chan, 2007b)
RO-SO ₃	~846		(Bondy et al., 2018)
S-O (-SO ₃ group)	1065		(Bondy et al., 2018)

1467