

- 2 the onset of the Late Pliensbachian Event in the Mochras
- 3 Borehole, Cardigan Bay Basin, NW Wales.'





SI Fig. 1: Clay mineral abundance diagram of the studied interval (951 – 918 mbs). Illite and smectite are most
 abundant, followed by kaolinite. Chlorite and I-S R1 are only abundant in the top part of the here studied record.





18 SI Fig. 2: The raw percentage of terrestrial phytoclast compared to the 3, 5, 10 and 15 step moving average.

19 Minor fluctuations of the terrestrial phytoclasts are observed on a metre scale around 30 %, however, the studied

20 interval contains 4 phases of higher proportions of terrestrial phytoclasts (marked in light grey shading). The

21 green graph represents the raw terrestrial phytoclast data and the black lines the 3, 5, 10 and 15 window moving

- average in this order.



SI Fig. 3: Palynofacies of the study interval 951 – 918 mbs. The particulate organic fraction of the samples
 studied is dominated by amorphous organic matter (AOM), followed by unstructured phytoclasts. Structured

37 phytoclasts and charcoal particles are relatively common. Palynomorphs, marine and terrestrial, are relatively

- sparse in all samples.





54 SI Fig. 4: The overall pattern of the microcharcoal abundance remains the same after correction for detrital
55 influx into the marine realm. Raw microcharcoal abundance (n/10g) compared to microcharcoal abundance
56 corrected for terrestrial influx with the XRF-elemental ratios (following Daniau *et al.*, 2013 and Hollaar *et al.*,
57 2021).

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70 SI Fig. 5: Simplified overview of the different terrestrial influx at times of relatively enhanced physical erosion 71 vs chemical weathering in the Cardigan Bay Basin observed over 405 kyr eccentricity cycles. A) shows the 72 scenario of relatively enhanced physical erosion on land, likely this led to a higher relative input of silt to sand 73 sized sediments and clay minerals illite and chlorite, which are indicative of physical erosion in the NW Tethys 74 region in the Early Jurassic (Merriman, 2006; Deconinck et al., 2019). I-S R1 is derived from the chemical 75 weathering of illite (Deconinck et al., 2019). Periods of less fine fraction sedimentations have been linked to 76 periods of lower carbonate dilution and/or enhanced preservation (Deconinck et al., 2019). B) illustrates the 77 scenario of relatively enhanced chemical weathering in the Cardigan Bay Basin. Thicker soil profiles likely 78 developed during this time, from both smectite and kaolinite. Elements associated with clays, such as Rb and Al, 79 were more abundant. Enhanced clay transport to the marine was associated with TOC-rich deposits in the 80 Cardigan Bay region (Deconinck et al., 2019). Limited burial diagenesis occurred in the Mochras core 81 (Deconinck et al., 2019). 82

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