



Chemical  
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R. Zhang et al.

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# Chemical characterization and source apportionment of PM<sub>2.5</sub> in Beijing: seasonal perspective

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Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



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**Chemical  
characterization and  
source  
apportionment**

R. Zhang et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## Abstract

In this study, 121 daily PM<sub>2.5</sub> (aerosol particle with aerodynamic diameter less than 2.5 μm) samples were collected from an urban site in Beijing in four months between April 2009 and January 2010 representing the four seasons. The samples were determined for various compositions, including elements, ions, and organic/elemental carbon. Various approaches, such as chemical mass balance, positive matrix factorization (PMF), trajectory clustering, and potential source contribution function (PSCF), were employed for characterizing aerosol speciation, identifying likely sources, and apportioning contributions from each likely source. Our results have shown distinctive seasonalities for various aerosol speciations associated with PM<sub>2.5</sub> in Beijing. Soil dust waxes in the spring and wanes in the summer. Regarding the secondary aerosol components, inorganic and organic species may behave in different manners. The former preferentially forms in the hot and humid summer via photochemical reactions, although their precursor gases, such as SO<sub>2</sub> and NO<sub>x</sub>, are emitted much more in winter. The latter seems to favorably form in the cold and dry winter. Synoptic meteorological and climate conditions can overwhelm the emission pattern in the formation of secondary aerosols. The PMF model identified six main sources: soil dust, coal combustion, biomass burning, traffic and waste incineration emission, industrial pollution, and secondary inorganic aerosol. Each of these sources has an annual mean contribution of 16, 14, 13, 3, 28, and 26 %, respectively, to PM<sub>2.5</sub>. However, the relative contributions of these identified sources significantly vary with changing seasons. The results of trajectory clustering and the PSCF method demonstrated that regional sources could be crucial contributors to PM pollution in Beijing. In conclusion, we have unraveled some complex aspects of the pollution sources and formation processes of PM<sub>2.5</sub> in Beijing. To our knowledge, this study is the first systematical study that comprehensively explores the chemical characterizations and source apportionments of PM<sub>2.5</sub> aerosol speciation in Beijing by applying multiple approaches based on a completely seasonal perspective.

### Chemical characterization and source apportionment

R. Zhang et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## 1 Introduction

Particulate matter (PM) is composed of various chemical components (Seinfeld, 1989). PM profoundly affects our living environments in terms of air quality (in close relation to public health), visibility, direct and indirect radiative forcing, climate effects and ecosystems (Watson, 2003; Streets et al., 2006; Andreae and Rosenfeld, 2008; Mahowald, 2011). Numerous epidemiological studies have demonstrated that longterm exposure to pronounced PM<sub>2.5</sub> increases morbidity and mortality (Dockery and Pope, 1995; Pope et al., 1995; Schwartz et al., 1996). Given its tiny size, fine mode PM (i.e. PM<sub>2.5</sub>, PM with aerodynamic diameter less than 2.5 μm) can readily penetrate the human bronchus and lung (Pope et al., 1995; Oberdorster, 2001). Through absorption and scattering of solar radiation and serving as cloud condensation nuclei, PM<sub>2.5</sub> extensively affects the global climate (Bardouki et al., 2003), and thus, the hydrological cycle (Ramanathan and Feng, 2009). The diverse effects of PM<sub>2.5</sub> could be a function of its complex chemical components and composition (He et al., 2009; Niwa et al., 2007; Malm et al., 2005; Eatough et al., 2006).

Due to the rapid economic and industrial developments and urbanization in the past few decades, there is an escalating increase in energy consumption and the number of motor vehicles in China, where air pollution has become ubiquitous (Chan and Yao, 2008). According to Shao et al. (2006), nearly 70 % of urban areas in China do not meet China's national ambient air quality standards, which are even much laxer than the air quality exposure standards/guidelines of the World Health Organization (WHO, 2005). The Beijing–Tianjin–Hebei region, the Yangtze River Delta, and the Pearl River Delta are of special concern because of their severe PM pollution, which can be explicitly shown by the spatial distribution of aerosol optical depth (AOD) retrieved by satellites (He et al., 2009; Lee et al., 2010). Three megacities that are representatives of each region, namely, Beijing, Shanghai, and Guangzhou, are the foci, because of their dense population. Coal is the primary energy source in China, and its consumption reached up to 1528 Mtce in 2005, accounting for nearly 70 % of the total energy consumption

### Chemical characterization and source apportionment

R. Zhang et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion







dust, crustal soil, coal combustion, secondary sulfate, secondary nitrate, biomass burning with municipal incineration, and vehicle emissions. All these studies were limited to a particular season and based on selected PM species.

To attain a better understanding of the chemical characteristics and sources of fine aerosols on a seasonal basis, we conducted a delicate investigation in Beijing. We continuously collected daily PM<sub>2.5</sub> samples at an urban site for four months, each of which in the respective seasons (i.e. spring, summer, autumn and winter). The samples were subjected to chemical measurements of various aerosol compositions as a whole, such as a suite of crustal and anthropogenic elements, major water-soluble ions, and OC/EC. Furthermore, we identified and apportioned the main sources to PM<sub>2.5</sub> by employing chemical mass closure construction and the PMF model in conjunction with trajectory cluster and potential source contribution function analyses according to the hybrid singleparticle Lagrangian integrated trajectory (HYSPLIT) model. This study will elucidate the source profile of PM<sub>2.5</sub> in different seasons and the relative contribution from each source in the complex urban airshed in Beijing and provide vital information in formulating the future air management framework to address the current alarming level of PM pollution in China which has been affecting the air quality on a vast regional scale.

## 2 Methodology

### 2.1 Sample and chemical analysis

#### 2.1.1 Sampling site

Beijing is located on the northern edge of the North China Plain, surrounded by the Yanshan Mountains in the west, north, and northeast (Fig. 1). According to the spatial distribution of fine AOD ranging from 0.0 to 1.0 that has been retrieved from Moderate-resolution Imaging Spectrometer (MODIS) sensors on board Terra and Aqua satellites

## Chemical characterization and source apportionment

R. Zhang et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

**Chemical  
characterization and  
source  
apportionment**

R. Zhang et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

(Fig. 1), Beijing is one of the PM<sub>2.5</sub> hot spots in China. The four seasons are characterized by variable meteorological conditions: spring by highspeed winds and low rainfall, summer by high temperature and frequent rain usually accounting for 75 % of annual rainfall, autumn by sunny days and northwest winds, and winter by cold and dry air. The population is 17.55 million. In 2009, the number of motor vehicles increased to 4.019 million, and energy consumption was equivalent to 65.73 million tons of standard coal (Beijing Statistics Yearbook, 2010). The sampling station was set up at the roof of the Science Building in Peking University (116.30° E, 39.99° N) 26 m a.g.l. A few field experimental campaigns have been conducted at this urban site (He et al., 2010; Guo et al., 2012). This site is located within the educational, commercial, and residential districts, and no main pollution sources exist nearby. Thus, the observations could be typical of the general urban pollution in Beijing.

### 2.1.2 Sample collection

Daily PM<sub>2.5</sub> samples were collected in April, July, and October 2009 and January 2010, representing spring, summer, autumn, and winter, respectively. Two collocated aerosol samplers (frmOMNI<sup>TM</sup>, BGI, USA) were used to collect PM<sub>2.5</sub> samples from 10 a.m. to 10 a.m. the next day simultaneously. The two substrates used in each sampler were 47 mm quartz filter (Whatman QM/A, England) and Teflon filter (pore size = 2 μm; Whatman PTFE, England). The flow rate was set at 5 L min<sup>-1</sup>. The quartz filters were baked at 800 °C for 3 h before use. The filter samples were stored at -18 °C until pretreatment.

### 2.1.3 Gravimetric weighing

Before and after each sampling, the PTFE filters were conditioned at 22 ± 1 °C in relative humidity of 35 ± 2 % for 24 h and then weighed in a weighing room by using an electronic balance with a detection limit of 1 μg (Sartorius, Gottingen, Germany). The corresponding PM<sub>2.5</sub> mass concentration of each filter was equal to the weight difference before and after sampling divided by the sampled air volume.

## 2.1.4 Chemical analysis of trace elements and water-soluble ions

Prior to extraction and digestion, each aerosol-laden PTFE membrane filter was cut into two equal halves with ceramic scissors. One half was subjected to Milli-Q water extraction for ionic measurement and the other half to acid digestion for elemental measurement. For the acid digestion, the polypropylene support O-ring on half of each PTFE filter sample was carefully removed with a ceramic knife from contamination. The filter samples were digested with an acid mixture (5 mL HNO<sub>3</sub> + 2 mL HF) by using an ultra-high throughput microwave digestion system (MARSXpress, CEM, Matthews, NC). A blank reagent and two filter blanks were prepared in each run following the same procedure used for the samples. All the acids used in this study were of ultra-pure grade (Merck, Germany). The detailed digestion method has been published elsewhere (Hsu et al., 2008). Another half of all filter samples were used for extraction with 20 mL Milli-Q purified water (specific resistivity = 18.2 MΩ cm; Millipore, Massachusetts, USA) for 1 h. The detailed extraction procedures have been described in Hsu et al. (2007, 2010a).

Ionic species (Na<sup>+</sup>, NH<sub>4</sub><sup>+</sup>, K<sup>+</sup>, Mg<sup>2+</sup>, Ca<sup>2+</sup>, F<sup>-</sup>, Cl<sup>-</sup>, SO<sub>4</sub><sup>2-</sup> and NO<sub>3</sub><sup>-</sup>) in the leachate were analyzed through a Dionex model ICS-1100 (for anions) and ICS-900 (for cations) ion chromatograph equipped with a conductivity detector (ASRS-ULTRA). Trace elements in the digestion solutions, including Al, Fe, Na, Mg, K, Ca, Ba, Ti, Mn, Co, Ni, Cu, Zn, Mo, Cd, Sn, Sr, Sb, Pb, Tl, Se, Ge, Cs, Ga, V, Cr, As, Se, and Rb, were analyzed by inductively coupled plasma–mass spectrometry (ICP–MS). Quality assurance and control of the ICP–MS was guaranteed by the analysis of a certified reference standard, NIST SRM-1648 (urban particulates). The resulting recoveries fell within ±10% of the certified values for most elements, except for Se, As, Cs, Sb, and Rb (±15%) (Hsu et al., 2009, 2010a).

## 2.1.5 OC and EC measurements

A punch of 0.526 cm<sup>2</sup> from each quartz filter was heated stepwise by a thermal/optical carbon analyzer (DRI 2001, Atmoslytic, US) in a pure helium atmosphere

### Chemical characterization and source apportionment

R. Zhang et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



at 140 °C (OC1), 280 °C (OC2), 480 °C (OC3), and 580 °C (OC4), and then in 2% O<sub>2</sub>/98% He atmosphere at 580 °C (EC1), 740 °C (EC2), and 840 °C (EC3) to convert any particulate carbon on the filter to CO<sub>2</sub>. After catalyzed by MnO<sub>2</sub>, CO<sub>2</sub> was reduced to CH<sub>4</sub>, which was then directly measured. Mass concentrations of OC and EC were obtained according to the IMPROVE protocol (Chow et al., 2007), OC = OC1 + OC2 + OC3 + OC4 + OP; EC = EC1 + EC2 + EC3-OP, where OP is the optical pyrolyzed OC. Detailed descriptions can be found in Zhang et al. (2012a).

## 2.2 Data analysis methods

### 2.2.1 Chemical mass closure

In this study, we constructed chemical mass closure (CMC) on a seasonal basis by considering mineral dust, SO<sub>4</sub><sup>2-</sup>, NO<sub>3</sub><sup>-</sup>, NH<sub>4</sub><sup>+</sup>, EC, particulate organic matter (POM), chloride salt (instead of sea salt; reason given below), trace element oxide (TEO), and biomass burning-derived K<sup>+</sup>. SO<sub>4</sub><sup>2-</sup>, NO<sub>3</sub><sup>-</sup>, and NH<sub>4</sub><sup>+</sup> can be regarded as the secondary inorganic aerosols.

The aluminosilicate (i.e. soil, dust, or mineral) component is often estimated through the following formula (Malm et al., 1994; Chow et al., 1994), which includes Si.

$$[\text{Mineral}] = 2.20\text{Al} + 2.49\text{Si} + 1.63\text{Ca} + 2.42\text{Fe} + 1.94\text{Ti}$$

$$[\text{Mineral}] = 1.89\text{Al} + 2.14\text{Si} + 1.40\text{Ca} + 1.43\text{Fe}$$

However, Si is volatilized as SiF<sub>4</sub> in the acid digestion of aerosol samples when using HF. Therefore, a few studies estimated Si from Al in the calculation of the mineral component (Hueglin et al., 2005). However, once Al is used in estimating Si concentrations, it generates relatively large uncertainty in the mineral component proportion as Si/Al mass ratios could largely vary in China's dust (Yan et al., 2012, and references therein). Accordingly, we adopted a straightforward method conventionally used

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

in estimating dust aerosols from Al:

$$[\text{Mineral}] = \text{Al}/0.07$$

where 0.07 is the average Al content (7%) reported by Zhang et al. (2003b). A similar estimation has been applied previously (Ho et al., 2006; Hsu et al., 2010b).

In estimating POM, we adopted a factor of 1.6 in converting OC to POM (Viidanoja et al., 2002), whereas a wide range of 1.4–2.2 has been utilized in previous investigations (Turpin and Lim, 2001; Andreae et al., 2008). The main determinants in selecting a conversion factor are the origin and age of the organic aerosols. The factor of 1.6 was employed in this study because the latest result shows a OM/OC ratio averaged at  $1.59 \pm 0.18$  in  $\text{PM}_{2.5}$  over China (Xing et al., 2013). This factor was used for the  $\text{PM}_{2.5}$  of Beijing by Dan et al. (2004), who also observed a similar seasonality for EC and OC and a OC/EC ratio (2–3) close to our results.

Sea salt is usually calculated as  $[\text{Sea salt}] = 1.82 \times \text{Cl}^-$  or  $= 2.54 \text{Na}^+$ . Given that Beijing is about 150 km away from East China's coastal oceans (i.e. Bohai Sea), sea spray-generated sea salt particles are not readily transported and are therefore insignificant to fine aerosols in Beijing. Nevertheless, dust blowing from Northern and Northwestern China is often associated with NaCl and  $\text{Na}_2\text{SO}_4$  from salt lake sediments and saline soils (Zhang et al., 2009). On the other hand, Cl may be essentially contributed by coal combustion in Beijing, particularly in winter (Yao et al., 2002). Thus, we considered chloride salt, instead of sea salt, as an individual component of  $\text{PM}_{2.5}$  aerosols in Beijing:  $[\text{Cl salt}] = [\text{Cl}^-] + [\text{Na}^+] + [\text{ss-Mg}^{2+}]$ . By considering chloride depletion in sea salt particles within the marine boundary layer because of the heterogeneous reaction, Hsu et al. (2010a) successfully evaluated such formula.

Following Landis et al. (2001), we estimated the contribution of heavy metals as metal oxides by employing the following equation:

$$\text{TEO} = 1.3 \cdot [0.5 \cdot (\text{Sr} + \text{Ba} + \text{Mn} + \text{Co} + \text{Rb} + \text{Ni} + \text{V}) + 1.0 \cdot (\text{Cu} + \text{Zn} + \text{Mo} + \text{Cd} + \text{Sn} + \text{Sb} + \text{Tl} + \text{Pb} + \text{As} + \text{Se} + \text{Ge} + \text{Cs} + \text{Ga})].$$

Chemical  
characterization and  
source  
apportionment

R. Zhang et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## Chemical characterization and source apportionment

R. Zhang et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

The enrichment factor (EF) of a given element (E) was calculated by using the formula  $EF = (E/Al)_{\text{Aerosol}} / (E/Al)_{\text{Crust}}$  (Hsu et al., 2010a), where  $(E/Al)_{\text{Aerosol}}$  is the ratio of the element to the Al mass in aerosols and  $(E/Al)_{\text{Crust}}$  is the ratio in the average crust (Taylor, 1964). The result of the EF is shown in Fig. S1. Elements with EFs of  $\leq 1.0$ , such as Cr and Y, were not considered, as they are of exclusive crustal origin. Elements with EFs between 1 and 5 were multiplied by a factor of 0.5, as they are possibly originated from two sources (i.e. anthropogenic and crustal sources). Elements with EFs  $\geq 5.0$  were multiplied by unity, as they are dominated by anthropogenic origins. Furthermore, the multiplicative factor was set at 1.3 so that metal abundance could be converted to oxide abundance, similar to those used by Landis et al. (2001). We also considered biomass burning-derived  $K^+$  ( $K_{\text{BB}}$ ) as an individual component, although  $K_{\text{BB}}$  salt may exist in the chemical forms of KCl and  $K_2SO_4$  (Pósfai et al., 2004), where both  $Cl^-$  and  $SO_4^{2-}$  have already been considered in other components.

### 2.2.2 PMF model

PMF is an effective source apportionment receptor model that does not require the source profiles prior to analysis and has no limitation on source numbers (Hopke, 2003; Shen et al., 2010). The principles of PMF can be found elsewhere in detail (Han et al., 2006; Song et al., 2006, Yu et al., 2013). In the present study, PMF 3.0 was employed with the inclusion of 34 chemical species in the model computation:  $PM_{2.5}$ , Al, Fe, Na, Mg, K, Ca, Ba, Ti, Mn, Co, Ni, Cu, Zn, Mo, Cd, Sn, Sb, Pb, V, Cr, As, Se, Rb,  $Na^+$ ,  $NH_4^+$ ,  $K^+$ ,  $Mg^{2+}$ ,  $Ca^{2+}$ ,  $Cl^-$ ,  $SO_4^{2-}$ ,  $NO_3^-$ , OC, and EC. Six physically realistic sources were identified.

### 2.2.3 Air mass back trajectory cluster

We calculated 48 h air mass back trajectories arriving at the sampling site ( $116.30^\circ E$ ,  $39.99^\circ N$ ) during our sampling period by using the National Oceanic and Atmospheric Administration (NOAA) HYSPLIT-4 model with a  $1^\circ \times 1^\circ$  latitude–longitude grid and the

final meteorological database. The six hourly final archive data were generated from the National Center for Environmental Prediction's Global Data Assimilation System (GDAS) wind field reanalysis. GDAS uses a spectral medium range forecast model. More details about the HYSPLIT model can be found at <http://www.arl.noaa.gov/ready/open/hysplit4.html> (NOAA Air Resources Laboratory). The model was run four times per day at starting times of 04:00, 10:00, 16:00, and 22:00 UTC (12:00, 18:00, 00:00, and 06:00 LT, respectively). The arrival level was set at 100 m above ground level. The method used in trajectory clustering was based on the GIS-based software TrajStat (<http://www.meteothinker.com/TrajStatProduct.aspx>).

#### 2.2.4 Potential source contribution function

The potential source contribution function (PSCF) is a method for identifying regional sources based on the HYSPLIT model. The zone of concern is divided into  $i \times j$  small equal grid cells. The PSCF value in the  $ij$ th cell is defined as  $m_{ij}/n_{ij}$ , where  $n_{ij}$  is designated as the numbers of endpoints that fall in the  $ij$ th cell and  $m_{ij}$  denotes the numbers of "polluted" trajectory endpoints in the  $ij$ th cell. In this analysis, average concentrations were treated as the "polluted" standard (Hsu et al., 2003). To better reflect the uncertainty in cells with small  $n_{ij}$  values (Polissar et al., 1999), the weighting function  $w_{ij}$  was adopted:

$$w_{ij} = \begin{cases} 1.00 & 80 < n_{ij} \\ 0.70 & 20 < n_{ij} \leq 80 \\ 0.42 & 10 < n_{ij} \leq 20 \\ 0.05 & n_{ij} \leq 10 \end{cases}$$

The study domain was in the range of 75°–130° E, 30°–60° N. The resolution was 0.5° × 0.5°.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

### 3 Results

#### 3.1 Annual average

Table 1 provides a statistical summary of the obtained data on atmospheric concentrations for PM<sub>2.5</sub>, Al (a tracer of aluminosilicate dust), water-soluble ions, OC, and EC during the sampling period. The annual mean PM<sub>2.5</sub> concentration reached 135 ± 63 μg m<sup>-3</sup>. This mean value is nearly three times higher than that (35 μg m<sup>-3</sup>) of the interim target-1 standard for annual mean PM<sub>2.5</sub> recommended by the WHO. The level of PM<sub>2.5</sub> in Beijing is much higher than other megacities around the world. In comparison with that of domestic cities, PM<sub>2.5</sub> seems to display a spatial tendency, increasing northward and decreasing southward (Zhang et al., 2012b). Such a spatial pattern may be related to the low rainfall and high dust in northern China (Qian et al., 2002, 2005). According to Wang et al. (2008), PM<sub>2.5</sub> concentrations in winter were much higher than in summer in 2001 to 2002. However, such a trend seemed to be reversed in 2005 to 2006, with rather higher concentrations in summer.

For the ionic concentrations, SO<sub>4</sub><sup>2-</sup> ranked the highest among the water-soluble ions analyzed, with an annual mean of 13.6 ± 12.4 μg m<sup>-3</sup>, followed by NO<sub>3</sub><sup>-</sup> (11.3 ± 10.8 μg m<sup>-3</sup>), NH<sub>4</sub><sup>+</sup> (6.9 ± 7.1 μg m<sup>-3</sup>), Ca<sup>2+</sup> (1.6 ± 1.4 μg m<sup>-3</sup>), Cl<sup>-</sup> (1.4 ± 2.2 μg m<sup>-3</sup>), K<sup>+</sup> (0.92 ± 0.75 μg m<sup>-3</sup>), Na<sup>+</sup> (0.46 ± 0.55 μg m<sup>-3</sup>), and Mg<sup>2+</sup> (0.16 ± 0.13 μg m<sup>-3</sup>). Such levels of mean concentrations are rather comparable with those measured in many Chinese cities such as Shanghai, Tianjin, Jinan, and Guangzhou (Yao et al., 2002; Tao et al., 2009; Gao et al., 2011; Gu et al., 2011). On average, the combination of SO<sub>4</sub><sup>2-</sup>, NO<sub>3</sub><sup>-</sup>, and NH<sub>4</sub><sup>+</sup>, which could be regarded as secondary inorganic aerosols, constituted the majority (88 %) of the total ionic concentrations, consistent with earlier studies (Yao et al., 2002; Duan et al., 2003). The annual mean concentrations of OC and EC reached up to 17.0 ± 10.0 and 5.0 ± 4.4 μg m<sup>-3</sup>, respectively. Such levels are close to those observed for regional sites across China (16.1 ± 5.2 μg m<sup>-3</sup> for OC and 3.6 ± 0.93 μg m<sup>-3</sup> for EC) by Zhang et al. (2008), who extensively measured

**Chemical  
characterization and  
source  
apportionment**

R. Zhang et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

carbonaceous aerosols around China; however, such concentrations are approximately half of those observed at urban sites ( $33.1 \pm 9.6 \mu\text{g m}^{-3}$  for OC and  $11.2 \pm 2.0 \mu\text{g m}^{-3}$  for EC) by Zhang et al. (2008).

### 3.2 Seasonality

As illustrated in Fig. 2, the seasonalities of  $\text{PM}_{2.5}$  and these primary species were characterized by distinctive features. The seasonality of  $\text{PM}_{2.5}$  concentration was not very evident and typical, with nearly equal concentrations of around  $140 \mu\text{g m}^{-3}$  in summer, autumn, and winter and a relative minimum ( $\sim 125 \mu\text{g m}^{-3}$ ) in spring. The minimum concentration typically occurs in summer because precipitation in Beijing is usually concentrated at that period. However, this is not the case, because the maximum concentrations of secondary sulfate and ammonium were observed in summer, arising from strong photochemistry and accounting for a large proportion ( $\sim 25\%$ ) of  $\text{PM}_{2.5}$  in Beijing (Yao et al., 2003).

In contrast to  $\text{PM}_{2.5}$ , sulfate and ammonium revealed a typical seasonality with higher concentrations in spring and summer and lower concentrations in autumn and winter, consistent with the seasonal variability of AOT (Xia et al., 2006). The summertime maximum concentrations of sulfate and ammonium were  $24$  and  $12 \mu\text{g m}^{-3}$ , respectively, which were higher than those in Beijing before 2003 ( $\sim 15$  and  $\leq 10 \mu\text{g m}^{-3}$ , respectively) (He et al., 2001; Duan et al., 2006; Wang et al., 2005) but rather comparable to those observed in the last few years (Okuda et al., 2011; Song et al., 2012). By contrast, the wintertime concentration of sulfate ( $8.5 \mu\text{g m}^{-3}$ ) was significantly reduced compared with earlier literature data (He et al., 2001; Hu et al., 2002; Wang et al., 2005). The decrease in wintertime sulfate concentration seemed to result from the effective control of  $\text{SO}_2$  emissions from coal combustion (Hao et al., 2005). High summertime sulfate concentration is ascribed to enhanced photochemistry during summer, and relatively high humidity accelerates the conversion rate of  $\text{SO}_2$  to the particulate form (Yao et al., 2003). However, the precursor  $\text{SO}_2$  concentrations are much higher in winter (Fig. S2) because of domestic heating (Hu et al., 2002b). In the present study,

## Chemical characterization and source apportionment

R. Zhang et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



**Chemical  
characterization and  
source  
apportionment**

R. Zhang et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

artificial biases, particularly of nitrate and ammonium, possibly occurred during sampling because no denuder and/or back-up filter was used to trap ammonia and nitric acid (Pathak et al., 2004). The maximum concentration ( $15.5 \mu\text{g m}^{-3}$ ) of nitrate was observed in spring rather than summer, which was different from that of sulfate (Wang et al., 2008). This observation may be ascribed to the volatility of ammonium nitrate, which is one of the main chemical forms of nitrate associations revealed by the ionic relationships. Thus, ammonium nitrate could evaporate at relatively high temperature. Besides, there are distinct emission sources for their respective precursor gases,  $\text{SO}_2$  and  $\text{NO}_x$ . The minimum concentration of nitrate was observed in winter. Nitrate levels could be a function of various factors in terms of emissions, such as exhaust, coal combustion, and biomass burning, and complex chemical processes with respect to photochemistry, heterogeneous reaction, renoxification, and gas–aerosol equilibrium. Rapid increases in on-road vehicular numbers in Beijing may offset any control measures for  $\text{NO}_x$  emissions, especially from coal combustion.

Crustally derived ions and elements, such as  $\text{Mg}^{2+}$ ,  $\text{Ca}^{2+}$ , and Al, waxed in the spring and waned in the summer, followed by significant increases toward autumn and winter. Such levels of seasonal mean concentrations and seasonalities are consistent with those observed in previous studies (Duan et al., 2006; Wang et al., 2005), which are related to dust storms and anthropogenic and fugitive dust. The seasonal concentrations of  $\text{Na}^+$  and  $\text{Cl}^-$  peaked in winter, consistent with Hu et al. (2002) and Wang et al. (2005). The seasonality of  $\text{Mg}^{2+}$  was distinguishable from that of  $\text{Na}^+$ , demonstrating the difference in their dominant sources.  $\text{K}^+$  had relatively higher concentrations in both spring and autumn than in summer and winter, which was closely associated with the agricultural burning around Beijing (Zheng et al., 2005). Such seasonality was distinct from those of previous studies, in which winter often had the highest concentration (He et al., 2001; Duan et al., 2006; Wang et al., 2005).

Both OC and EC had similar seasonal patterns of waxing in winter and waning in spring (for EC) or summer (for OC). Zhang et al. (2008) observed a persistently common seasonality for both OC and EC at 18 background, regional, and urban stations

## Chemical characterization and source apportionment

R. Zhang et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

in China, namely, a maximum in winter and a minimum in summer. The seasonalities may be governed by the variability in emission strengths and meteorology. For instance, lower-molecular weight semi-volatile organic compounds are mostly in gaseous phase at high temperature in summer (Yassaa et al., 2001). The OC/EC mass ratio 2.0 indicates the presence of secondary organic matter (Chow et al., 1996). In this work, the OC/EC ratios mostly fell within the range of 2–5, with mean ratios of 4.8, 2.5, 3.0, and 2.7 in spring, summer, autumn, and winter, respectively (Fig. 3). These figures are very similar to those observed in previous studies (Duan et al., 2006; Zhang et al., 2008), which suggests the relative domination of secondary organic aerosols in spring but of primary sources in other seasons (Zhang et al., 2008). Another reason for the relatively higher springtime OC/EC mass ratio may be the open biomass burning source, consistent with higher  $K^+$  in the spring (Fig. 2), as the aerosols from open biomass burning are generally characterized by elevated OC/EC ratios (Cao et al., 2007).

### 3.3 Stoichiometric analyses of cations and anions

Note that equivalent concentrations ( $\mu\text{eq m}^{-3}$ ) are used throughout this section. Figure 4 shows the scatter plots of (a)  $\text{Mg}^{2+}$  vs.  $\text{Na}^+$ , (b)  $\text{Mg}^{2+}$  vs.  $\text{Ca}^{2+}$ , (c)  $\text{Cl}^-$  vs.  $\text{Na}^+$ , and (d)  $\text{Cl}^-$  vs.  $\text{K}^+$ . Figure 5 shows the scatter plots of (a) total cations vs. total anions, (b)  $\text{NH}_4^+$  vs.  $\text{SO}_4^{2-}$ , (c)  $\text{NH}_4^+$  vs.  $[\text{SO}_4^{2-} + \text{NO}_3^-]$ , (d)  $[\text{NH}_4^+ + \text{Ca}^{2+}]$  vs.  $[\text{SO}_4^{2-} + \text{NO}_3^-]$ , (e)  $[\text{NH}_4^+ + \text{Ca}^{2+} + \text{Mg}^{2+}]$  vs.  $[\text{SO}_4^{2-} + \text{NO}_3^-]$ , and (f)  $[\text{NH}_4^+ + \text{Ca}^{2+} + \text{Mg}^{2+}]$  vs.  $[\text{SO}_4^{2-} + \text{NO}_3^- + \text{Ex-Cl}^-]$ . Ex- $\text{Cl}^-$  is the excessive  $\text{Cl}^-$ , defined as the excessive amount of  $\text{Cl}^-$  relative to the amount sea salt can sustain with  $\text{Na}^+$  as the tracer of sea salt:  $\text{Ex-Cl}^- = \text{Cl}^- - [\text{Na}^+] \cdot 1.17$ , where 1.17 is the typical  $\text{Cl}/\text{Na}^+$  equivalent ratio of average seawater (Chester, 1990). If the resulting Ex- $\text{Cl}^-$  is negative, then no  $\text{Cl}^-$  excess exists. In other words,  $\text{Cl}^-$  is totally contributed by sea salt and is even depleted by heterogeneous reactions. Nevertheless, total  $\text{Na}^+$  does not necessarily originate from sea salt alone, but could partially come from dust. The resultant biases are hence likely insignificant.

**Chemical  
characterization and  
source  
apportionment**

R. Zhang et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

Figure 4a illustrates that  $\text{Mg}^{2+}$  mostly comes from non-sea salt sources, except in wintertime, because the regression slopes that represent the  $\text{Mg}^{2+}/\text{Na}^+$  ratios (1.56, 0.70, 0.68, and 0.25 for spring, summer, autumn, and winter, respectively) are clearly deviated from the ratio (0.23) of average seawater (Chester, 1990). Instead, the dominant source of  $\text{Mg}^{2+}$  is mineral dust, mainly carbonate minerals (Li et al., 2007), as reflected by the good correlations (0.97, 0.83, 0.90, and 0.80 for spring, summer, autumn, and winter, respectively) between  $\text{Mg}^{2+}$  and  $\text{Ca}^{2+}$  (Fig. 4b). Similarly, most  $\text{Cl}^-/\text{Na}^+$  ratios (1.70, 1.21, 2.51, and 2.25 for spring, summer, autumn, and winter, respectively) in  $\text{PM}_{2.5}$  are larger than the mean ratio (1.17) of seawater, except the summertime samples (Fig. 4c). This difference indicates the dominance of the non-sea salt sources, of which the most likely contributor of  $\text{Cl}^-$  is coal combustion (Yao et al., 2002), particularly in winter when  $\text{Cl}^-$  and  $\text{SO}_2$  are maximal (Figs. 2 and S2). In summer, air masses are dominated by the southerly monsoon from Bohai (as supported by the trajectories below), leading to a mean  $\text{Cl}^-/\text{Na}^+$  ratio close to that of average seawater. Moreover, the correlations between  $\text{K}^+$  and  $\text{Cl}^-$  largely varied with the seasons, with better correlation and higher ratios in autumn and winter and moderate correlations and lower ratios (less than unity) in summer and spring (Fig. 4d). Thus, the results suggested that  $\text{K}^+$  was not present in chemical form  $\text{KCl}$  at high temperature, but as  $\text{K}_2\text{SO}_4$  (Pósfai et al., 2004). By contrast, low temperature in winter may favor the presence of  $\text{KCl}$ .

Furthermore, the ratio of total cation concentration to total anion concentration is averaged at near unity throughout the year (Fig. 5a), which indicates excellent charge balance in  $\text{PM}_{2.5}$  and high data quality. Figure 5b shows good correlations between  $\text{NH}_4^+$  and  $\text{SO}_4^{2-}$  for the annual data set, with ratios (represented by the slope of the linear regression line) between 1.25 and 1.77 (all higher than unity). These good correlations reveal the dominance of  $(\text{NH}_4)_2\text{SO}_4$  (Ianniello et al., 2011), rather than  $\text{NH}_4\text{HSO}_4$  and the possible full neutralization of  $\text{SO}_4^{2-}$  by  $\text{NH}_4^+$  throughout the year. We further considered the combination of  $\text{NO}_3^-$  and  $\text{SO}_4^{2-}$  in this charge balance analysis (Fig. 5c). The resulting  $\text{NH}_4^+$  to  $[\text{SO}_4^{2-} + \text{NO}_3^-]$  ratios ranged from 0.83 to 0.94 (all lower than unity),

**Chemical  
characterization and  
source  
apportionment**

R. Zhang et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

which demonstrates the presence of  $\text{NH}_4\text{NO}_3$  in the fine mode aerosols. Moreover, ratios lower than unity suggest that nitrate may be present in other chemical forms than  $\text{NH}_4\text{NO}_3$ . Heterogeneous reactions between  $\text{NO}_x$  (and its products, such as  $\text{HNO}_3$  and  $\text{N}_2\text{O}_5$ ) and dust carbonate are often observed in northern China (Li and Shao, 2009).

5 Accordingly, we examined the correlations of  $[\text{NH}_4^+ + \text{Ca}^{2+}]$  vs.  $[\text{SO}_4^{2-} + \text{NO}_3^-]$  (Fig. 5d) and of  $[\text{NH}_4^+ + \text{Ca}^{2+} + \text{Mg}^{2+}]$  vs.  $[\text{SO}_4^{2-} + \text{NO}_3^-]$  (Fig. 5e), given that the good correlations between  $\text{Mg}^{2+}$  and  $\text{Ca}^{2+}$  suggest the possible existence of water-soluble Mg in reacted carbonate dust. These ions are strongly correlated throughout the year, with high coefficients (all 0.99 or higher) and slopes of regression lines around unity. These

10 correlations indicate that nitrate is partly present in  $\text{Ca}(\text{NO}_3)_2$  and  $\text{Mg}(\text{NO}_3)_2$ , not just in  $\text{NH}_4\text{NO}_3$ . We assumed that fine sized sulfate is exclusively associated with ammonium and that  $\text{Na}^+$  is present only in the associated NaCl. However, these assumptions may not always be true because  $\text{Na}_2\text{SO}_4$  is observed in dust particles from dried lakes in northern China (Zhang et al., 2009a) and NaCl could react with nitric acid to form

15  $\text{NaNO}_3$  via heterogeneous reaction (Hsu et al., 2007). Therefore, based on the aforementioned equivalent interrelationships and assumptions, we quantitatively estimated that the former two chemical forms ( $\text{Ca}(\text{NO}_3)_2$  and  $\text{Mg}(\text{NO}_3)_2$ ) represent  $\sim 20\%$  of the total nitrate and that  $\text{NH}_4\text{NO}_3$  is the dominant association accounting for the remaining  $\sim 80\%$ . The elevated  $\text{Cl}^-/\text{Na}^+$  ratio ( $> 1.17$ ) shows that excessive  $\text{Cl}^-$  seemed to be

20 attributed to coal combustion rather than sea salt particles from dried salt lake sediment. The addition of excessive  $\text{Cl}^-$  (Fig. 5f) insignificantly changed the correlations of positive and negative charges. Nevertheless, we noted that in wintertime, the equivalent ratio improved from 1.10 to 1.00, which indicates the presence of chloride salts such as KCl,  $\text{CaCl}_2$ , and  $\text{MgCl}_2$  other than NaCl at low ambient temperature (Ianniello et al., 2011). KCl may have originated from biomass burning, and  $\text{CaCl}_2$  and  $\text{MgCl}_2$

25 could have been formed through heterogeneous reactions between the dust carbonate and HCl emitted from coal combustion.

### 3.4 Chemical mass closure

By employing the methods in Sect. 2.2.1, we constructed the CMC of  $PM_{2.5}$  in Beijing on a seasonal and annual basis. The reconstructed  $PM_{2.5}$  mass concentrations were compared with the gravimetric  $PM_{2.5}$  mass concentrations, as shown in Fig. 6, which shows a good correlation with one another in each season and throughout the year. However, the ratios seasonally changed, with higher ratios of 0.82 in spring and 0.75 in winter and lower ratios of 0.59 in summer and 0.68 in autumn. The proportions of all specific components in  $PM_{2.5}$  together with the unidentified constituents as a whole are schematically illustrated by five pie charts for the four seasonal and annual cases (Fig. 7). Overall, the major components are secondary inorganic aerosols (combination of sulfate, nitrate, and ammonium), mineral dust, and POM, which account for each ~20%, albeit with seasonal variations. The minor components include EC, chloride salt, potassium salt, and TEO, each of which represents less than 5%. Specifically, the proportions of mineral dust are maximal (33.4%) in spring, minimal (only 8.2%) in summer, and intermediate (23.6 and 28.9%) in the other two seasons, consistent with the tendency of seasonal Al concentrations. The totals of secondary inorganic species ( $SO_4^{2-}$ ,  $NO_3^-$ , and  $NH_4^+$ ) have the largest proportion (27% to 30%) in spring and summer and a minimal percentage (< 15%) in autumn and winter. However, sulfate peaks were noted in summer (15.4%), whereas nitrate peaks were observed in spring (11.1%). Ammonium decreased from around 5 to 7% in spring and summer to half (2 to 3%) in autumn and winter. The POM fractions largely varied as follows: summer (14.7%) < spring (18.9%) < autumn (22.1%) < winter (28.6%). EC and chloride salt exhibited the largest proportions (5.2 and 3.3%, respectively) in winter. Potassium salt and TEO had slightly higher proportions in spring and autumn than in summer and winter.

Both the primary and secondary components of  $PM_{2.5}$  in Beijing are equally important, albeit with seasonal variability, which is typical of PM pollution in China (Shao et al., 2006). In general, given that the seasonal variability in  $PM_{2.5}$  mass

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

concentrations is relatively small, temporal trends in the proportions of each component of  $PM_{2.5}$  resemble the atmospheric concentrations of their corresponding chemical species. The likely factors for such seasonalities are partially addressed in Sect. 3.2 and discussed in detail in the following two sections.

5 On average, the unidentified components reached 28.6% of the total  $PM_{2.5}$ . They also showed seasonal variability, with the smallest (15.9%) in spring when dust was prevalent, and the largest (42.5%) in summer when secondary inorganic aerosol formation was favorable. Such high uncertainties in the CMCs were caused by the water absorption of water-soluble components in the weighing environment, though relative humidity was controlled (Speer et al., 1997; Tsai and Kuo, 2005). The absorption likely led to positive biases in  $PM_{2.5}$  concentrations. Alternatively, such uncertainties may be partly due to the volatilization of  $NH_4NO_3$  and organic matter, particularly in summer and autumn during the storage of the weighted samples prior to extraction, which may have resulted in negative biases in the specific components. Another likely reason for the non-match of the reconstructed and gravimetric mass concentrations is the varying factors used in transferring a given analyzed species (e.g. OC and Al) to a certain component (e.g. POM and mineral soils) (Rees et al., 2004; Hsu et al., 2010a; Yan et al., 2012). For example, a few studies adopted a factor of only 1.4 for converting OC content to organic matter (Duan et al., 2006; Song et al., 2007; Guinot et al., 2007). Another study obtained a much higher POM/OC mean ratio over China (Xing et al., 2013) of up to  $1.92 \pm 0.39$  based on a mass balance method. If we adopt this higher ratio, the unidentified percentage would be reduced by 3%. In the present study, the  $EF_{crust}$  of Ca averages at 2.6, which shows its enrichment relative to average crust composition. In Beijing, fine sized Ca-rich dust is partly attributed to construction activities. Therefore, we may underestimate carbonate abundance in the mineral component estimated from Al concentration alone (Guinot et al., 2007). This may have resulted in the underestimation ( $\sim 2\%$ ) of the total mass reconstructed.

## Chemical characterization and source apportionment

R. Zhang et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion







species that could represent respective contribution sources (Fig. 2). For instance, we compared the maximal and minimal contributions of mineral dust in spring and summer, respectively, which are consistent with the seasonality of aerosol AI. Dust storms are essentially responsible for springtime dust aerosols, whereas in autumn and winter, fugitive dust from construction and the resuspension of street dust are the main contributors. Obviously, the reconstructed time series of daily concentrations from coal combustion reveals a pronounced wintertime maximum, consistent with those of aerosol  $\text{Cl}^-$  (Fig. 2) and even gaseous  $\text{SO}_2$  (Fig. S2). Moreover, the time series of biomass burning contributions show relatively higher concentrations in spring and autumn and lower concentrations in summer and winter, consistent with the seasonality of  $\text{K}^+$ . For traffic and waste burning emissions, the resulting time series do not reveal evident seasonality, corresponding with the seasonalities of nitrate and some trace metals, such as Pb, Cu, Sb, and Cd (Fig. S3). Industrial pollution has higher contributions in summer and autumn, possibly corresponding with the seasonalities of Zn and Cr. However, such seasonality is inconsistent with that of OC, with a wintertime maximum, because OC may be from various sources, including the former five sources identified. Coal combustion has the largest contribution in winter, and low temperature in winter facilitates the formation of secondary organic aerosols. SIA has higher contributions in summer and spring, mirroring the seasonalities of sulfate, nitrate, and ammonium. This result is definitely related to the photochemistry that accounts for SIA formation. The formed SIA species may not appear in their original emission sources (i.e. coal combustion, biomass burning, traffic exhausts, waste incineration, and industrial pollution), but in the SIA component. Based on the PMF results based and chemical data in January and August 2004, Song et al. (2007) found that the most predominant sources of  $\text{PM}_{2.5}$  are coal combustion in winter and secondary aerosols in summer, along with other significant sources, such as motor vehicle emissions, road dust, and biomass burning. The PMF-modeled results seem to be promising because the corresponding time series of each source's contribution are very consistent with the observations.

## Chemical characterization and source apportionment

R. Zhang et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## 4.2 Regional sources deduced from trajectory and PSCF analyses

The regional sources and transport of air pollutants exert a profound impact on local air quality in Beijing (e.g. Wang et al., 2004). To address this issue, both trajectory clustering and PSCF methods were employed. The 48 h back trajectories starting at 100 m from Beijing were computed by using the HYSPLIT model of NOAA (http://www.arl.noaa.gov/ready.html). Four clusters were made (Fig. 10): northwestern (including western, NW), northern (N), eastern (from northeastern to southeastern, E), and southern (S) directions. The NW cluster was further differentiated into two types, namely, fast ( $NW_f$ ) and slow ( $NW_s$ ), according to the motion speed ( $\leq 7 \text{ m sec}^{-1}$  for  $NW_s$  and  $> 7 \text{ m sec}^{-1}$  for  $NW_f$ ) and distance of air parcels. The classification is consistent with the spatial distribution of fine AOD retrieved by remote sensing (Fig. 1).

Table 3 summarizes the percentages of each trajectory cluster in the total on annual and seasonal bases and the corresponding mean concentrations of  $PM_{2.5}$  and various aerosol species associated within each trajectory cluster. Annually, the trajectory clusters are dominated by both NW and S, accounting for 44 % and 34 %, respectively. The E and N clusters represent the rest (15 % and 7 %, respectively). However, the variability is large and season-dependent. For instance, the predominant clusters are N (30 %) and S (44 %) in spring, S (73 %) in summer,  $NW_f$  (50 %) in autumn, and  $NW_f$  (88 %) in winter. The resulting mean concentrations of main aerosol species seasonally vary with certain types of air masses. In winter, a few  $PM_{2.5}$  pollution cases (only 12 % of the wintertime trajectories) with mean concentration as high as  $209 \mu\text{g m}^{-3}$  are associated within the E trajectories that passed over Hebei and Liaoning Provinces, where heavy industries are concentrated in certain cities (e.g. Tianjin, Tangshan, Dalian, Shenyang). However, in spring, summer, and autumn, high  $PM_{2.5}$  ( $> 150 \mu\text{g m}^{-3}$ ) is preferentially associated with the S trajectory cluster. Overall, the general patterns agree with the spatial distribution of the MODIS retrieved fine AOD around Beijing (Fig. 1).

Furthermore, we applied an alternative approach called PSCF to explore the likely regional sources and transport pathways of various  $PM_{2.5}$  associated speciations, such

### Chemical characterization and source apportionment

R. Zhang et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

**Chemical  
characterization and  
source  
apportionment**

R. Zhang et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

as sulfate, nitrate, ammonium, OC, EC, and mineral dust in Beijing, as illustrated in Fig. 11. A few main features were found: (a) sulfate, nitrate, and ammonium have similar spatial patterns, with higher values in the east to the south, covering Tianjing, Shijiazhuang, and Zhengzhou; (b) both OC and EC show similar spatial distribution, with higher values in the northwest, the south, and the northeast, covering the border of Hebei and Shanxi Provinces, Inner Mongolia, the border of the Hebei, Shanxi, and Henan Provinces, and the area from Tianjing to Shenyang; (c) the higher value for mineral aerosols is localized in the northwest and the south; and (d) for these six aerosol speciations, the southern area appears to be a common hot spot. The overall PSCF results are rather consistent with the spatial distribution of fine AOD (Fig. 1). The statistics obtained from the trajectory clustering (Table 3) shows that the southern air masses bring high levels of secondary inorganic and carbonaceous aerosols and the northwestern air masses are enriched in mineral dust and carbonaceous aerosols. Sun et al. (2006) and Street et al. (2007) found that the S sector has much higher secondary species, such as sulfate, nitrate, and ammonium. During haze-fog events in Beijing, chemical constituents of secondary inorganic aerosols are also much higher when the winds blow from the south. Such high amounts of secondary fine-sized aerosols in southern air parcels may be related to high humidity (water vapor) and enhanced heterogeneous reaction in clouds/fog, aside from strong photochemistry. The association of high dust with the NW trajectories is consistent with Wang et al. (2004) and Yu et al. (2011).

**4.3 Implications for atmospheric chemistry, PM control measures, and climate**

Rigorous efforts exerted for air pollutant governance prior to the 2008 Beijing Summer Olympics, such as changing the energy source structure, reducing local dust emissions, controlling vehicle exhaust emissions, and relocating major industrial emitters, have achieved air quality improvement during the games. Effective control of coarse PM pollution seemed possible and the main urban air pollutants became finer PM ( $PM_{2.5}$ ). The annual mean concentration of  $PM_{2.5}$  in Beijing is nearly three times and



implemented locally, regionally, and nationally in China. Such measures would improve pollution abatement, public health, and climatic modeling capacity.

## 5 Summary

The levels of daily  $\text{PM}_{2.5}$  concentrations are still elevated in Beijing throughout the year, with an annual mean of up to  $135 \pm 63 \mu\text{g m}^{-3}$ , which is several times higher than the 24 h exposure standard and guideline recommended by the WHO. Seasonality is not very evident, although the highest occurred in winter and the lowest in spring. Distinctive seasonalities occurred for various aerosol species. Sulfate and ammonium peaked in summer, mainly because of photochemistry, whereas the maximum of nitrate was observed in spring, which is attributed to the heterogeneous reaction between dust carbonate and nitric acid (the product of  $\text{NO}_x$ ) along with the photochemical process. Crustally derived species, such as  $\text{Ca}^{2+}$ ,  $\text{Mg}^{2+}$ , and Al, showed typical seasonality, waxing in spring and waning in summer. The maximum of OC, BC, and Cl was observed in winter, and their minimum in summer or spring, which suggests a relation with fossil fuel (mainly coal) combustion processes. The correlations between total cation and anion equivalent concentrations were excellent with a mean ratio around unity, which indicates good charge balance. The mean  $\text{NH}_4^+/\text{SO}_4^{2-}$  equivalent concentration ratios were higher than unity throughout the year, which reveals that  $\text{NH}_4^+$  could fully neutralize sulfuric acid. The correlations between  $\text{NH}_4^+$  and  $[\text{SO}_4^{2-} + \text{NO}_3^-]$  suggest that ammonium could partially be present as  $\text{NH}_4\text{NO}_3$ . Furthermore, the correlations between  $[\text{NH}_4^+ + \text{Ca}^{2+}]$  and  $[\text{SO}_4^{2-} + \text{NO}_3^-]$  demonstrate the presence of  $\text{Ca}(\text{NO}_3)_2$  and  $\text{NH}_4\text{NO}_3$ , which was due to the heterogeneous reaction between dust carbonate and nitric acid. A few chloride salts may form, particularly in winter when low temperature is favorable for their occurrence.

CMCs were successfully constructed on a seasonal basis, although with an unidentified fraction averaged at 28.6%. The major aerosol speciation considered in this study

ACPD

13, 9953–10007, 2013

### Chemical characterization and source apportionment

R. Zhang et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



emission strengths of primary aerosols and gaseous precursors, the dynamically variable synoptic weather conditions and circulation patterns also have a crucial role in governing the anomalies of PM (both fine and coarse size) concentrations.

**Supplementary material related to this article is available online at:**

**<http://www.atmos-chem-phys-discuss.net/13/9953/2013/acpd-13-9953-2013-supplement.pdf>.**

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ACPD

13, 9953–10007, 2013

## Chemical characterization and source apportionment

R. Zhang et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

## Chemical characterization and source apportionment

R. Zhang et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

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**Chemical  
characterization and  
source  
apportionment**

R. Zhang et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

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## Chemical characterization and source apportionment

R. Zhang et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

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R. Zhang et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

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**Chemical  
characterization and  
source  
apportionment**

R. Zhang et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

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## Chemical characterization and source apportionment

R. Zhang et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

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**Chemical  
characterization and  
source  
apportionment**

R. Zhang et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

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**Chemical  
characterization and  
source  
apportionment**

R. Zhang et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



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**Chemical  
characterization and  
source  
apportionment**

R. Zhang et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

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## Chemical characterization and source apportionment

R. Zhang et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

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## Chemical characterization and source apportionment

R. Zhang et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

**Table 1.** Statistical summary showing the means (with one standard deviation) and ranges of atmospheric concentrations for PM<sub>2.5</sub> (in  $\mu\text{g m}^{-3}$ ) and selected species (in  $\text{ng m}^{-3}$ ) in the entire sampling (annual) and four season (months) periods.

Species	Annual	Spring	Summer	Autumn	Winter
PM <sub>2.5</sub>	135 ± 63 39–355	126 ± 59 39–280	138 ± 48 41–226	135 ± 55 45–251	139 ± 86 48–355
SO <sub>4</sub> <sup>2-</sup>	13.6 ± 12.4 0.9–52.8	14.7 ± 11.5 2.3–52.8	23.5 ± 14.5 2.5–52.0	7.9 ± 7.4 0.9–25.7	8.5 ± 8.6 1.3–34.4
NO <sub>3</sub> <sup>-</sup>	11.3 ± 10.8 0.3–63.8	15.5 ± 13.7 1.3–63.8	11.8 ± 8.2 1.8–31.5	10.7 ± 11.0 0.3–34.7	7.3 ± 8.1 1.6–35.5
NH <sub>4</sub> <sup>+</sup>	6.9 ± 7.1 0.1–39.1	7.5 ± 8.1 0.6–39.1	11.0 ± 6.9 0.5–23.9	4.7 ± 5.8 0.1–17.7	4.5 ± 5.7 0.3–23.3
Cl <sup>-</sup>	1.42 ± 2.18 0.03–10.34	0.72 ± 0.81 0.04–3.74	0.30 ± 0.56 0.03–3.06	1.12 ± 0.98 0.09–3.71	3.52 ± 3.32 0.19–10.34
Na <sup>+</sup>	0.46 ± 0.55 0.04–2.82	0.31 ± 0.18 0.08–0.94	0.17 ± 0.09 0.04–0.42	0.30 ± 0.22 0.05–1.06	1.08 ± 0.80 0.11–2.82
K <sup>+</sup>	0.92 ± 0.75 0.03–3.66	1.08 ± 0.71 0.14–3.14	0.66 ± 0.47 0.20–2.47	1.13 ± 0.90 0.03–3.66	0.81 ± 0.77 0.05–2.53
Mg <sup>2+</sup>	0.16 ± 0.13 0.02–1.04	0.24 ± 0.20 0.03–1.04	0.07 ± 0.03 0.02–0.16	0.16 ± 0.07 0.06–0.31	0.18 ± 0.09 0.06–0.45
Ca <sup>2+</sup>	1.6 ± 1.5 0.2–11.3	2.6 ± 2.2 0.2–11.3	0.6 ± 0.3 0.2–1.7	1.7 ± 1.0 0.5–4.2	1.5 ± 0.9 0.5–4.0
Al	1.8 ± 1.5 0.1–6.9	2.5 ± 1.7 0.3–6.6	0.7 ± 0.4 0.1–2.0	2.0 ± 1.4 0.3–6.7	2.1 ± 1.5 0.7–6.9
OC	16.9 ± 10.0 5.9–58.6	13.7 ± 4.4 5.9–23.7	11.1 ± 1.8 7.4–16.6	17.8 ± 5.6 7.5–26.2	24.9 ± 15.6 8.5–58.6
EC	5.0 ± 4.4 0.6–28.1	2.8 ± 1.1 0.6–5.8	4.2 ± 1.2 1.5–6.8	5.3 ± 2.8 1.3–12.1	7.5 ± 7.4 1.2–28.1

## Chemical characterization and source apportionment

R. Zhang et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

**Table 2.** Relative contributions from six identified sources of PM<sub>2.5</sub> in Beijing within the one-year and four-season periods.

Source	Spring	Summer	Autumn	Winter	Annual
Soil dust	23 %	3 %	18 %	16 %	15 %
Coal combustion	5 %	1 %	7 %	57 %	18 %
Biomass burning	19 %	6 %	17 %	7 %	12 %
Traffic and waste incineration emission	5 %	4 %	4 %	2 %	4 %
Industrial pollution	14 %	32 %	42 %	12 %	25 %
SIA	34 %	54 %	13 %	6 %	26 %

## Chemical characterization and source apportionment

R. Zhang et al.

**Table 3.** Mean concentrations (in  $\mu\text{g m}^{-3}$ ) of  $\text{PM}_{2.5}$  and selected aerosol speciations in the identified trajectory clusters within the one-year (annual) and four season period. Also given are the percentages of each trajectory cluster classified in the one-year and four season periods. For details on trajectory clustering, please refer to the text.

Air-mass type	Annual					Spring				Summer			Autumn			Winter		
	NW <sub>s</sub>	NW <sub>f</sub>	N	E	S	NW <sub>s</sub>	N	E	S	NW <sub>s</sub>	N	E	S	NW <sub>f</sub>	E	S	NW <sub>f</sub>	E
Percent (%)	11.9	32.4	7.4	14.6	33.7	8.7	29.6	17.4	44.3	9.6	8.7	8.7	73.	49.6	26.8	23.6	88.2	11.8
PM <sub>2.5</sub>	148	111	87	110	172	145	70	108	167	108	110	63	155	113	144	173	131	209
Sulfate	10.2	5	6.3	10.9	25.4	11.9	4.8	13.6	23.1	11.1	9.6	7.6	28.3	4.8	10.9	11.2	7	19.5
Nitrate	11.2	5.1	4.3	10.2	19.2	14	4.8	14.4	24.4	4.5	4	3.5	14.3	6.4	13.4	16.9	6.3	14.9
Ammonium	5.4	2.3	2.3	5.7	13.2	4.7	1.7	8	12.2	4.8	3.7	3.4	13.4	2.4	6.5	7.6	3.7	10.9
OC	22.2	18.1	10.8	13.8	16.4	15.5	10	11.5	16.7	12.4	12.2	9.4	11	15.6	18.7	21.6	23.8	33.6
EC	6.5	4.9	2.5	3.7	5.4	3.7	1.9	2.2	3.4	3.5	3.6	2.6	4.5	4	5.9	7.5	7	11.4
Mineral	31.3	36.5	26	16.4	20.8	64.8	32.8	15.3	39.1	15	21	4.8	8.5	33.3	19.4	30.4	31.3	21.8

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

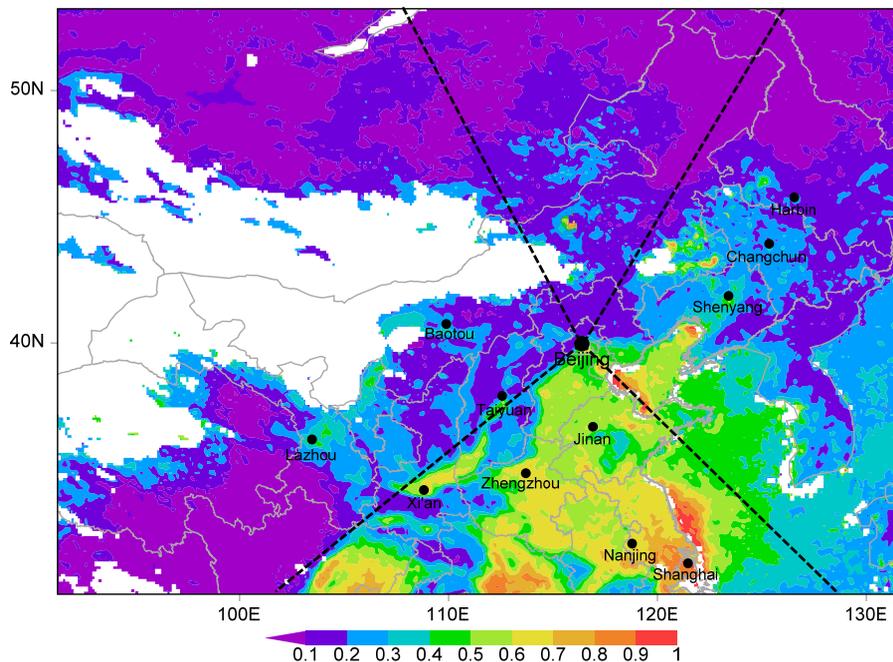
Printer-friendly Version

Interactive Discussion



## Chemical characterization and source apportionment

R. Zhang et al.

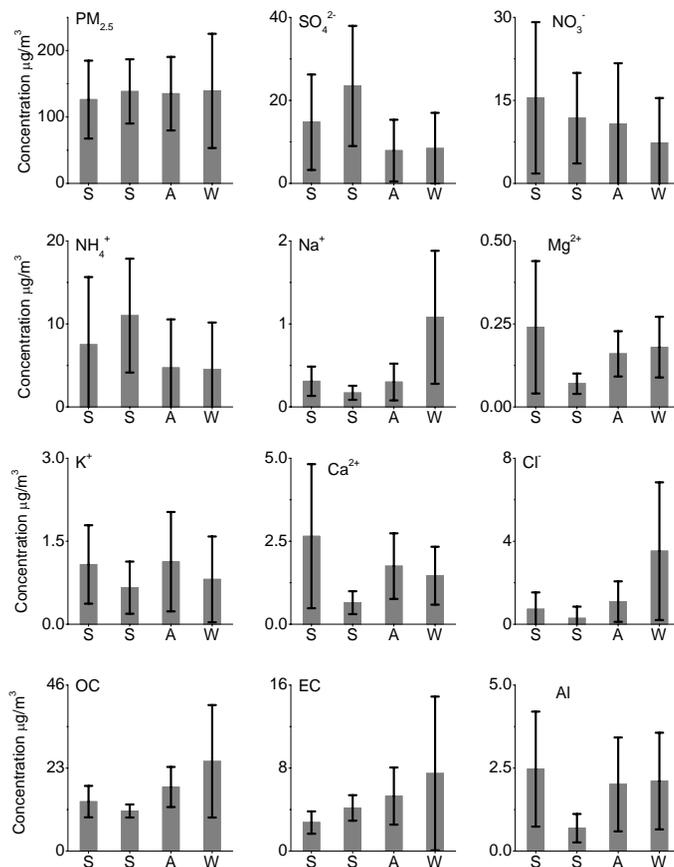


**Fig. 1.** Sampling location ( $116.30^{\circ}$  E,  $39.99^{\circ}$  N) on a regional map superimposed with spatial distribution of annual mean fine aerosol optical depth (AOD) retrieved from MODIS satellite remote sensing in 2009. Dashed lines define four regions based on the trajectory clustering results discussed in Sect. 4.2 (seen in the text). Also shown are several major cities around Beijing.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[⏪](#)[⏩](#)[⏴](#)[⏵](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

## Chemical characterization and source apportionment

R. Zhang et al.

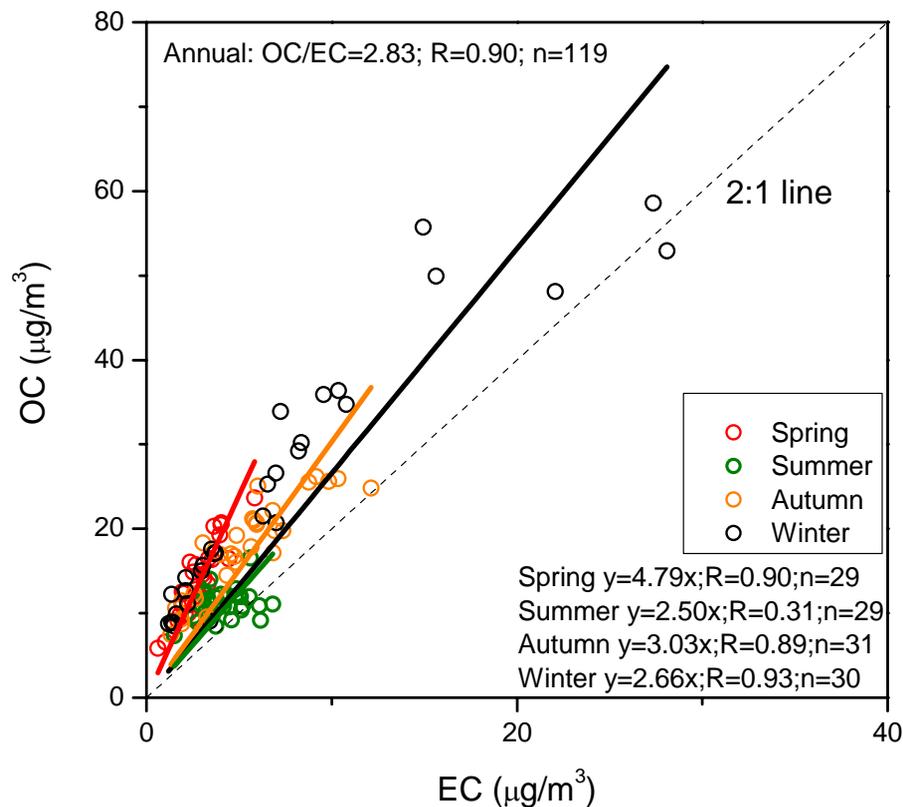


**Fig. 2.** Seasonal variations of PM<sub>2.5</sub> mass concentration and associated species, including SO<sub>4</sub><sup>2-</sup>, NO<sub>3</sub><sup>-</sup>, NH<sub>4</sub><sup>+</sup>, Na<sup>+</sup>, Mg<sup>2+</sup>, K<sup>+</sup>, Ca<sub>2</sub><sup>+</sup>, Cl<sup>-</sup>, OC, EC, and Al concentrations. Shown here are the mean and one standard deviation for each bar.

[Title Page](#)
[Abstract](#)
[Introduction](#)
[Conclusions](#)
[References](#)
[Tables](#)
[Figures](#)
[Back](#)
[Close](#)
[Full Screen / Esc](#)
[Printer-friendly Version](#)
[Interactive Discussion](#)

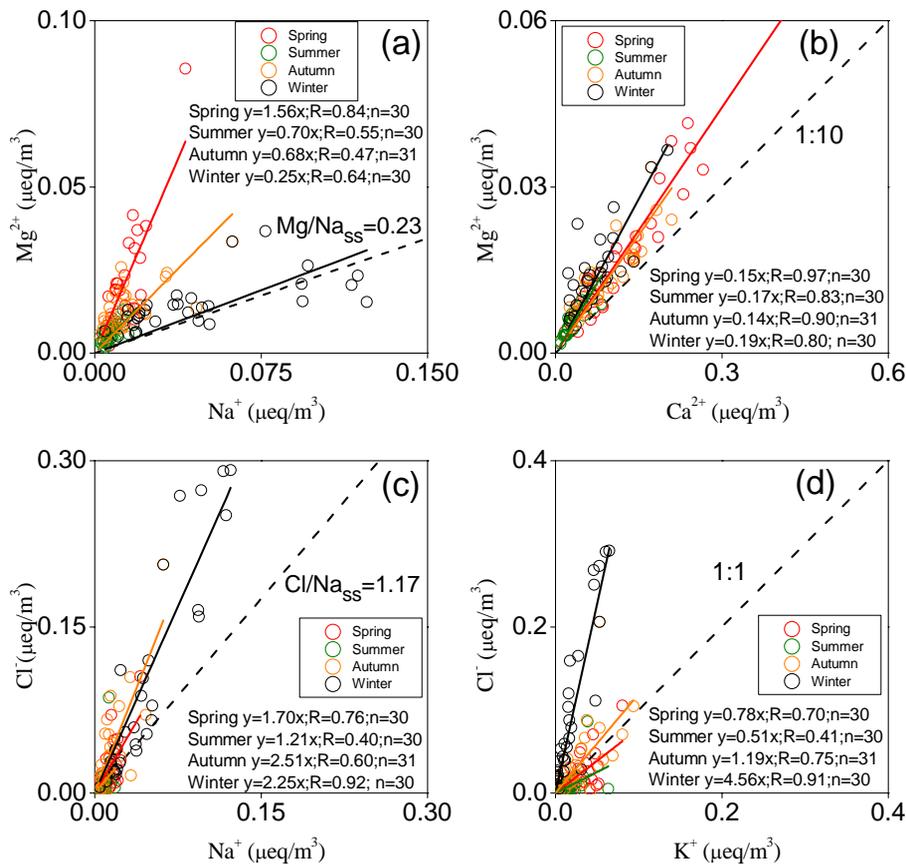
Chemical  
characterization and  
source  
apportionment

R. Zhang et al.



**Fig. 3.** Scatter plot showing the correlation between OC (y-axis) and EC (x-axis) in  $\text{PM}_{2.5}$  collected from Beijing. Different symbols denote the four seasons. Linear regression equations are given in the annual and seasonal cases.

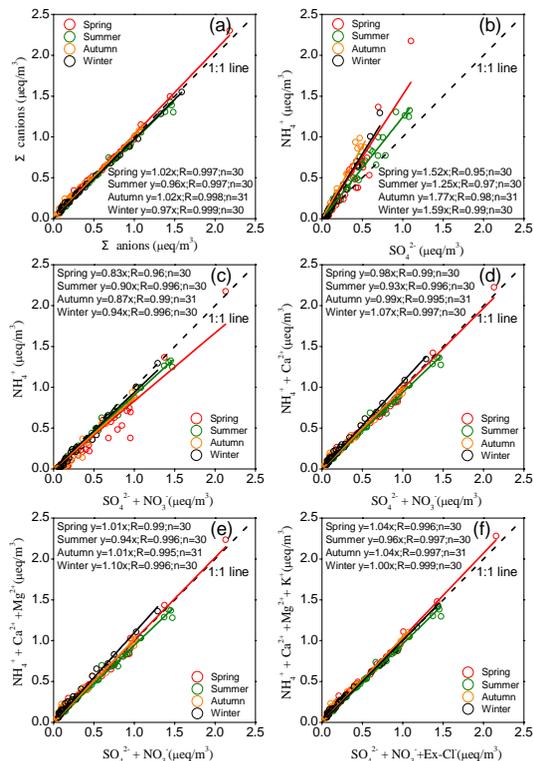
[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[⏪](#)[⏩](#)[⏴](#)[⏵](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)



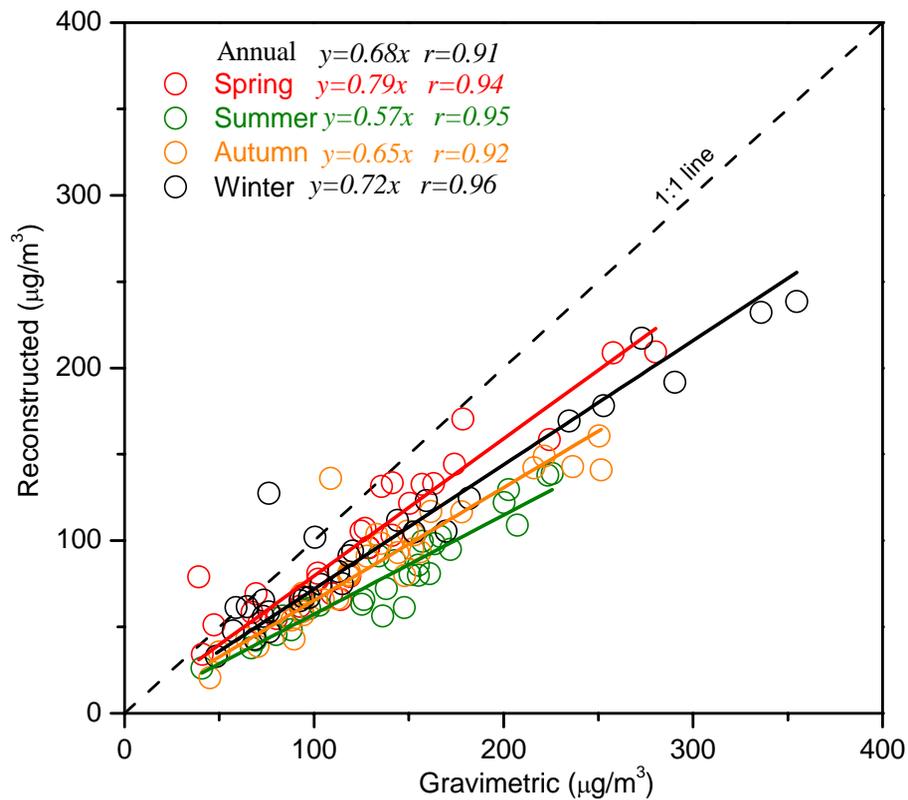
**Fig. 4.** Correlations between certain cations and anions: **(a)**  $Mg^{2+}$  versus  $Na^+$ , **(b)**  $Mg^{2+}$  versus  $Ca^{2+}$ , **(c)**  $Cl^-$  versus  $Na^+$ , and  $Cl^-$  versus  $K^+$ . Four seasons are considered and represented by different color symbols. Linear regression equations are also given. In panels **(a)** and **(c)**, the dashed lines indicate the  $Mg/Na$  and  $Cl/Na$  equivalent concentration ratios in seawater (sea salt), respectively.

Chemical  
characterization and  
source  
apportionment

R. Zhang et al.



**Fig. 5.** Same as Fig. 4, but for **(a)** total cations versus total anions, **(b)**  $\text{NH}_4^+$  versus  $\text{SO}_4^{2-}$ , **(c)**  $\text{NH}_4^+$  versus  $[\text{SO}_4^{2-} + \text{NO}_3^-]$ , **(d)**  $[\text{NH}_4^+ + \text{Ca}^{2+}]$  versus  $[\text{SO}_4^{2-} + \text{NO}_3^-]$ , **(e)**  $[\text{NH}_4^+ + \text{Ca}^{2+} + \text{Mg}^{2+}]$  versus  $[\text{SO}_4^{2-} + \text{NO}_3^-]$ , and **(f)**  $[\text{NH}_4^+ + \text{Ca}^{2+} + \text{Mg}^{2+}]$  versus  $[\text{SO}_4^{2-} + \text{NO}_3^- + \text{Ex-Cl}^-]$ . The detailed definition of Ex- $\text{Cl}^-$  (excessive- $\text{Cl}^-$ ) can be found in the text. The 1 : 1 lines are given for comparison.



**Fig. 6.** Scatter plot showing the correlations between the  $\text{PM}_{2.5}$  mass concentrations reconstructed from the chemical mass balance method and obtained from gravimetric measurement, which are presented on a seasonal basis. Linear regression lines are shown with equations for the four season period, along with the annual case.

**Chemical  
characterization and  
source  
apportionment**

R. Zhang et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

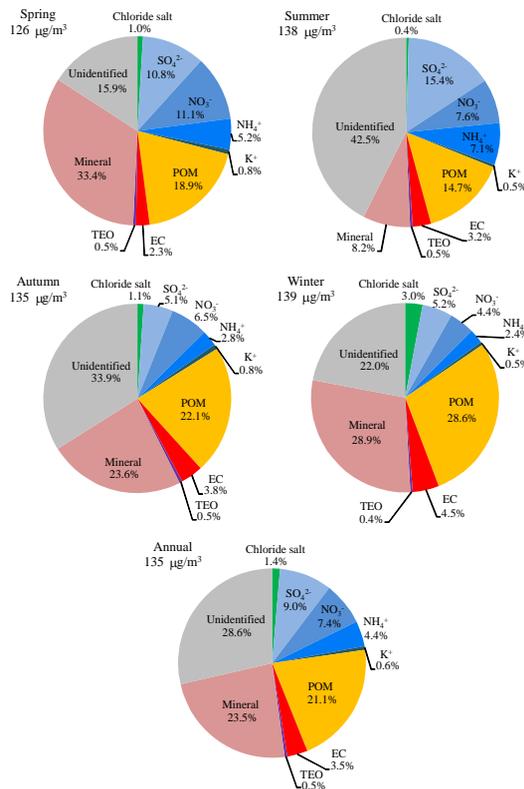
Printer-friendly Version

Interactive Discussion



## Chemical characterization and source apportionment

R. Zhang et al.



**Fig. 7.** Pie-charts showing the constructed chemical mass closures for PM<sub>2.5</sub> in Beijing: **(a)** spring; **(b)** summer; **(c)** autumn; **(d)** winter; and **(e)** annual. The components include mineral dust, secondary aerosol ions (sulfate, nitrate, and ammonium), POM, EC, trace element oxides (TEO), chloride salt, and biomass burning-derived potassium. Other than the identified components, unidentified fractions comprise a significant portion of PM<sub>2.5</sub>.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

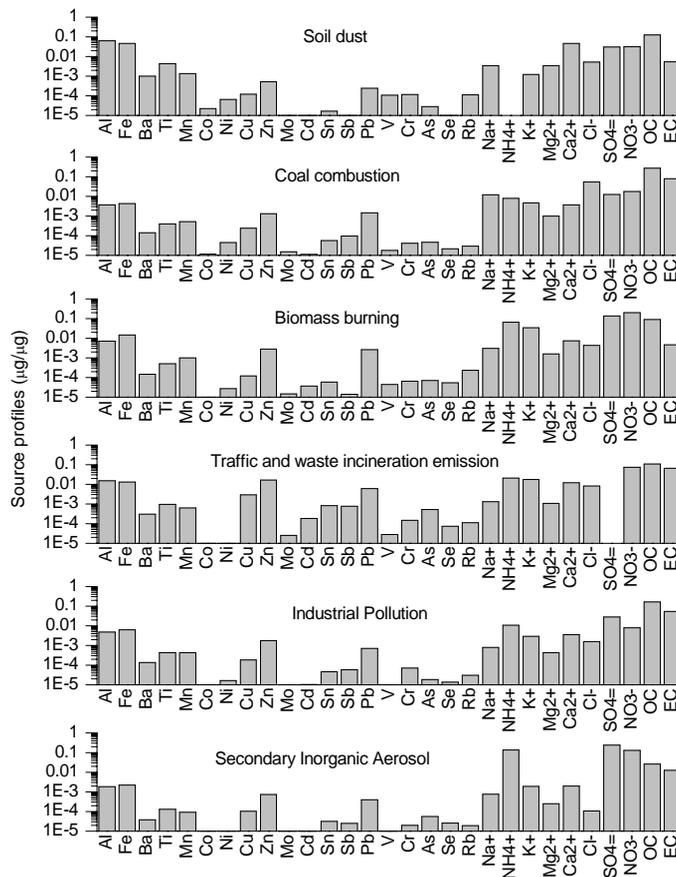
Full Screen / Esc

Printer-friendly Version

Interactive Discussion

## Chemical characterization and source apportionment

R. Zhang et al.



**Fig. 8.** Profiles of six sources identified from the PMF model, including soil dust, coal combustion, biomass burning, traffic and waste incineration emission, industrial pollution, and secondary inorganic aerosol (from the upper to the lower panels).

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

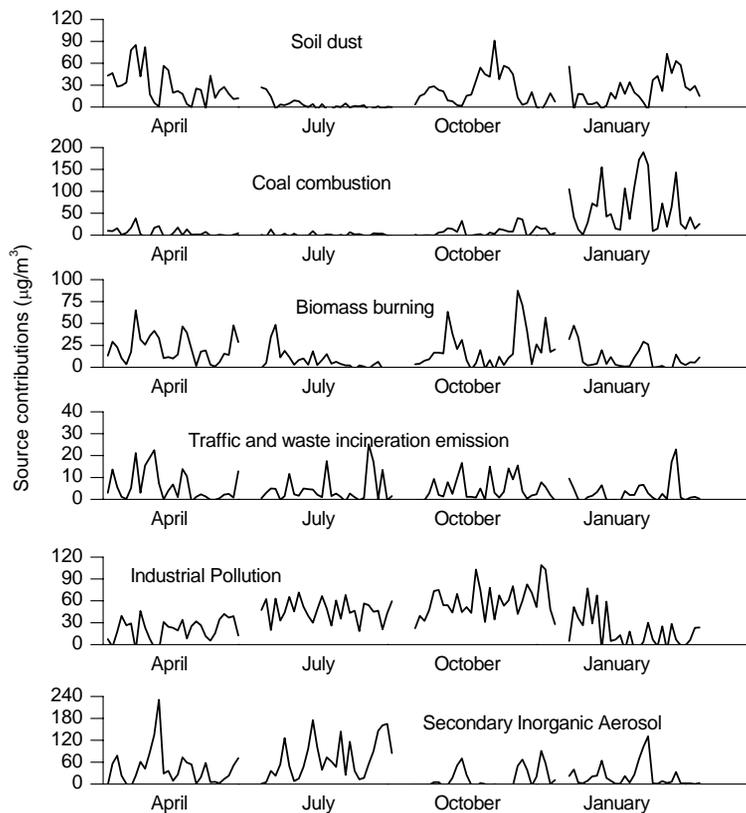
Full Screen / Esc

Printer-friendly Version

Interactive Discussion

Chemical  
characterization and  
source  
apportionment

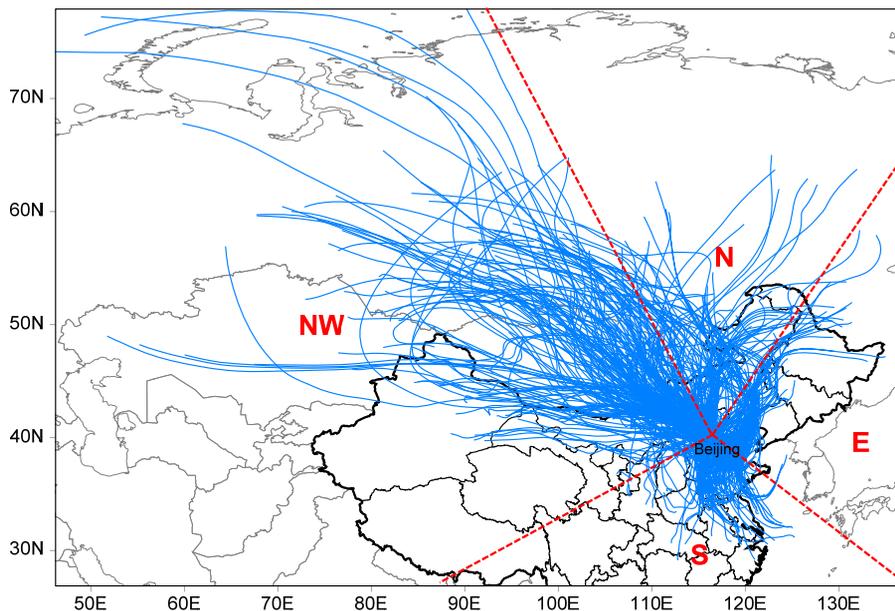
R. Zhang et al.



**Fig. 9.** Time-series of daily contributions from each identified source, including soil dust, coal combustion, biomass burning, traffic and waste incineration emission, industrial pollution, and secondary inorganic aerosol (from the upper to the lower panels) during the study period between April 2009 and January 2010.

**Chemical  
characterization and  
source  
apportionment**

R. Zhang et al.



**Fig. 10.** Analytical results of the 48 h air-mass back trajectories at 100 m elevation during the sampling periods, which were run four times per day. Four regions were defined based on the trajectory clustering results, namely, NW, N, E, and S regions.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[⏪](#)[⏩](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

Chemical  
characterization and  
source  
apportionment

R. Zhang et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

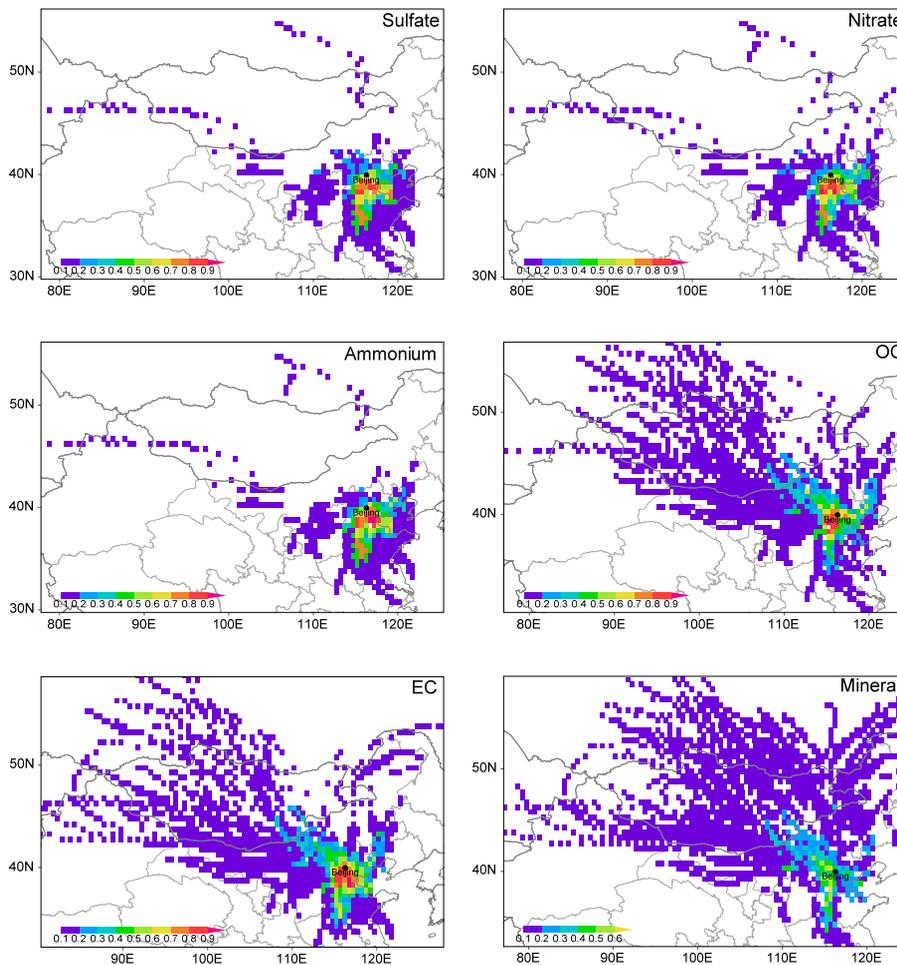
Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

**Fig. 11.** The PSCF maps for sulfate, nitrate, ammonium, OC, EC, and mineral.